

Incorporating Breaking Wave Predictions in Spectral Wave Forecast Models

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Motivational Aspects

- **wave breaking at sea** - widespread air-sea interfacial process with very significant geophysical and maritime importance
- present spectral wave forecast models **do not provide explicit forecasts** of breaking wave properties.
- recent advances in understanding wave breaking have made it possible to redress this deficiency
- to describe a novel methodology that adds to standard spectral wave model output - **accurate forecasts** of
 - (i) the **spectral density of breaking crest length per unit area** and
 - (ii) the associated **breaking strength**

We did this initially for the **dominant wind waves** and have now extended it across the **full spectrum**

St Jude's Storm shutdown: Eurostar and Monday morning train services cancelled across south as coast is lashed by 25ft waves

- Amendments and cancellations on First Capital Connect, Southeastern, Greater Anglia and Stansted Express
- Also disruption on East Coast, c2c, First Great Western, Southern, Gatwick Express and South West Trains
- Ferries from Poole and Weymouth to Guernsey & Jersey cancelled and hovercrafts to Isle of Wight suspended
- About 60 flights cancelled at London Heathrow Airport tomorrow but none yet at Gatwick, Stansted and Luton
- Forecasters warn houses face damage, trees falling and power cuts in biggest storm to hit Britain in a decade
- Wales and South West England will be hit first early tomorrow morning with winds of up to 90mph expected
- Boy, 14, believed to have drowned today after swimming with friends in waves off Newhaven in East Sussex
- Canoeist dies after being pulled from swollen River Tees near Barnard Castle, County Durham, after capsizing

By MARK DUELL, TARA BRADY and JONATHAN PETRE





Modeling Wave Breaking Spectrally

Radiative transfer equation (deep water, no currents)

The radiative transfer equation for describing the evolution of the wave height spectrum $F(k)$ is given by:

$$\frac{\partial F}{\partial t} + c_g \cdot \nabla F = S_{in} + S_{nl} + S_{diss}$$

where

- $F=F(k, \theta)$ is the directional wave spectrum
- c_g is the group velocity
- $S_{tot} = S_{in} + S_{nl} + S_{ds}$ is the total source term.
- S_{in} is the **atmospheric input** spectral source term
- S_{nl} is the **nonlinear spectral transfer** source term representing nonlinear wave-wave interactions within the spectrum
- S_{ds} is the spectral **dissipation rate** here taken as due primarily to wave breaking

Saturation Threshold-based Dissipation Rate S_{ds}

- based on treating spectral bands as nonlinear wave groups. Uses a low power of the spectral saturation ratio (~steepness ratio) to simulate observed threshold behaviour [*extension of Alves & Banner (JPO, 2003)*]

$$S_{ds}(k, \theta) = [C_1 * D * (\tilde{\sigma} - \tilde{\sigma}_T) / \tilde{\sigma}_T]^{a_1} + C_2 * D * E_{tot} * k_p^2] (\sigma / \sigma_m)^{a_2} \omega F(k, \theta)$$



‘local S_{ds} ’



‘non-local S_{ds} ’

This formulation uses

- normalized azimuthally-integrated saturation: $k^4 F(k) / \theta(k) = (2\pi)^4 f^5 F(f) / 2g^2$
- measured threshold of the normalized spectral saturation (Banner et al., JPO, 2002) with $a_1=2$
- tail exponent $a_2 = 4$ to match dissipation to input behavior in the spectral tail
- nonlocal dissipation rate component
- coefficient D for the local S_{ds} : non-dimensional and linear in the wind speed to match to the wind input term.
- C_1 and C_2 constants

Modified Jansen Wind Input

$$S_{in}(k, \theta) = \varepsilon \beta(k, \theta) \omega(u_*^{red}(k) \cos \theta / c)^2 * E(k, \theta)$$

$$\beta(k, \theta) = J_2 \mu (\ln(\mu))^4 / \kappa^2$$

where $J_2=1.6$ (Janssen (1991) used 1.2).

$$\beta(k, \theta) = 0 \quad \text{for} \quad \mu > 1$$

$$\mu(k, \theta) = (u_*/c)^2 (gz_0/u_*^2) \exp(J_1 \kappa / (u_* \cos \theta / c)^2)$$

$$z_0 = \frac{0.01 u_*^2}{g} / \sqrt{1 - C_0 (\tau_w / \tau)}$$

$$u_*^{red}(k_n) = \sqrt{[\tau_{tot} - J_0 \sum_{i=1}^n (\tau_w(i) + \tau_{bw}(i))] / \rho_{air}}$$

Brief description of the methodology

$\Lambda(\mathbf{c})$ is the spectral density of breaking wave crest length per unit area

$\Pi(\mathbf{c})$ is the spectral density of the total wave crest length per unit area

The breaking probability $P_{br}(\mathbf{c})$ for wave scales \mathbf{c} is defined as:

$$Pr_{br}(\mathbf{c}) = \frac{\int \mathbf{c} \Lambda d\mathbf{c}}{\int \mathbf{c} \Pi d\mathbf{c}} \quad \sim \quad \frac{\text{passage rate of breaking wave crests}}{\text{passage rate of wave crests}}$$

$\Lambda(\mathbf{c})$: spectral density of *breaking wave crest length* per unit area with velocities in the range $(\mathbf{c}, \mathbf{c}+d\mathbf{c})$ (Phillips,1985)

$$b \frac{\rho}{g} c^5 \Lambda(c) dc \quad \text{wave energy dissipation rate at scale } c$$

$$b \frac{\rho}{g} c^4 \Lambda(c) dc \quad \text{momentum flux from waves of scale } c \text{ to currents}$$

The sea state threshold variable used for breaking probability was the normalised spectral saturation

$$\tilde{\sigma}(k) = \sigma(k) / \langle \theta(k) \rangle$$

where $\sigma(k)$ is the azimuth-integrated spectral saturation given by

$$\sigma(k) = k^4 \Phi(k)$$

$$= (2\pi)^4 f^5 G(f) / 2g^2$$

and $\langle \theta(k) \rangle$ is the mean spectral spreading width given by

$$\langle \theta(k) \rangle = \int_{-\pi}^{\pi} (\theta - \bar{\theta}) F(k, \theta) k d\theta / \int_{-\pi}^{\pi} F(k, \theta) k d\theta$$

where $\bar{\theta}$ is the mean wave direction, and $\Phi(k)$, $G(f)$ and $F(k, \theta)$ are, respectively, the spectra of wave height as a function of scalar wavenumber, frequency and vector wavenumber.

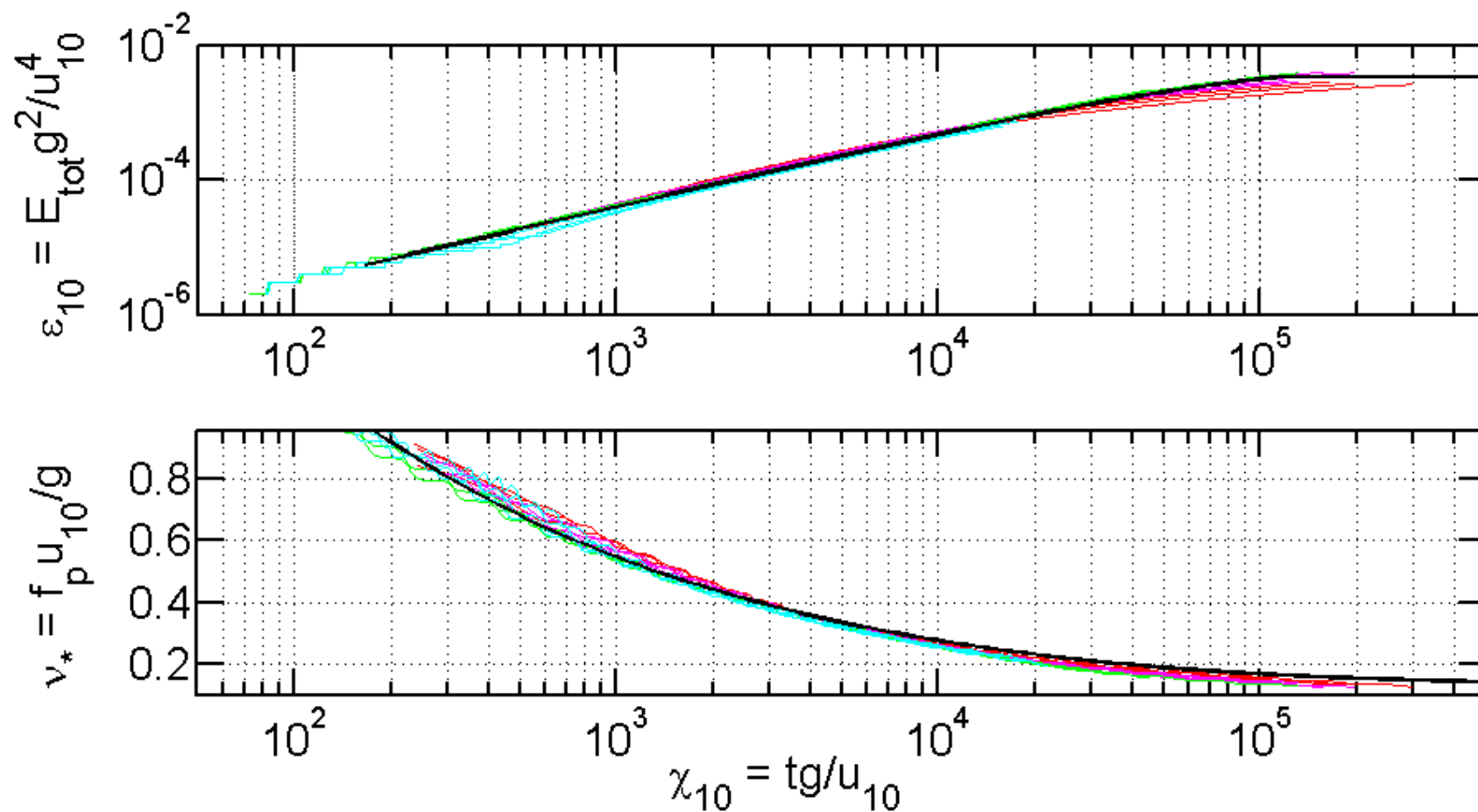
$$\text{Pr}_{\text{br}}(\tilde{\sigma}) = H(\tilde{\sigma} - \tilde{\sigma}_{\text{T}}) * \alpha_{\text{br}} * (\tilde{\sigma} - \tilde{\sigma}_{\text{T}})^{0.5}$$

$$\text{b}_{\text{br}}(\tilde{\sigma}) = H(\tilde{\sigma} - \tilde{\sigma}_{\text{T}}) * c_{\text{br}} * (\tilde{\sigma} - \tilde{\sigma}_{\text{T}})^{1.0}$$

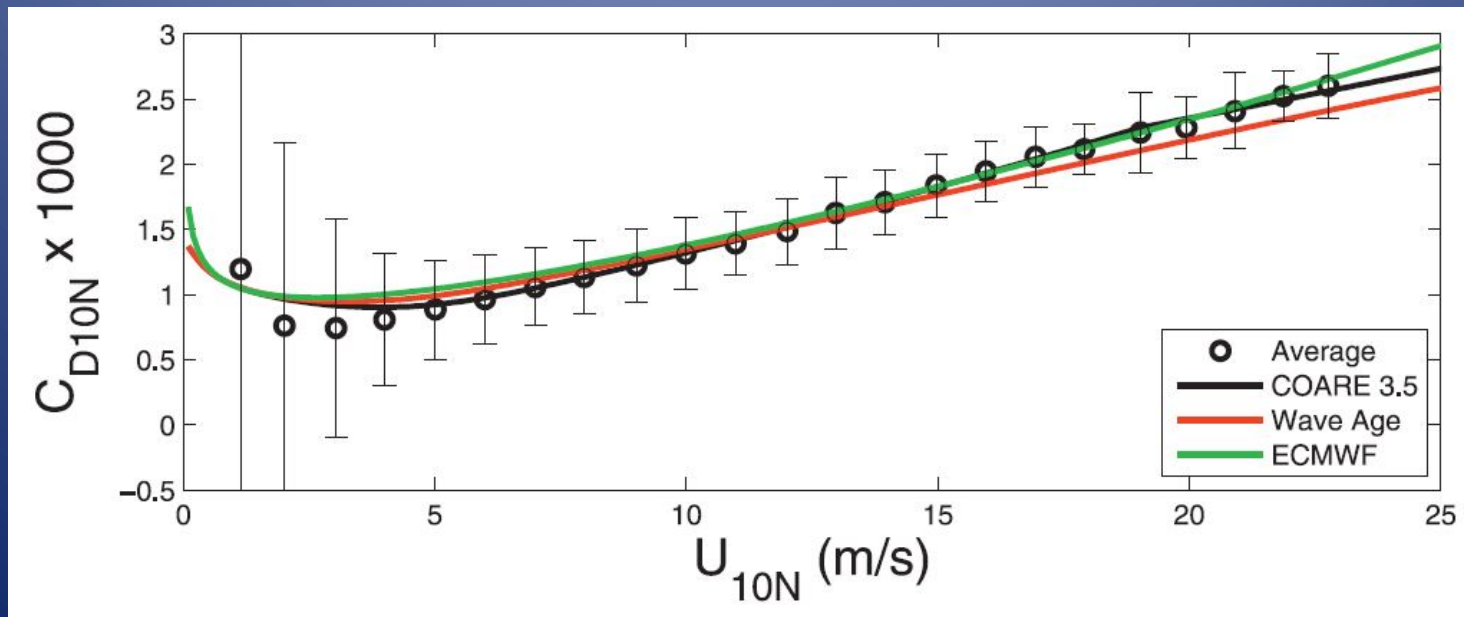
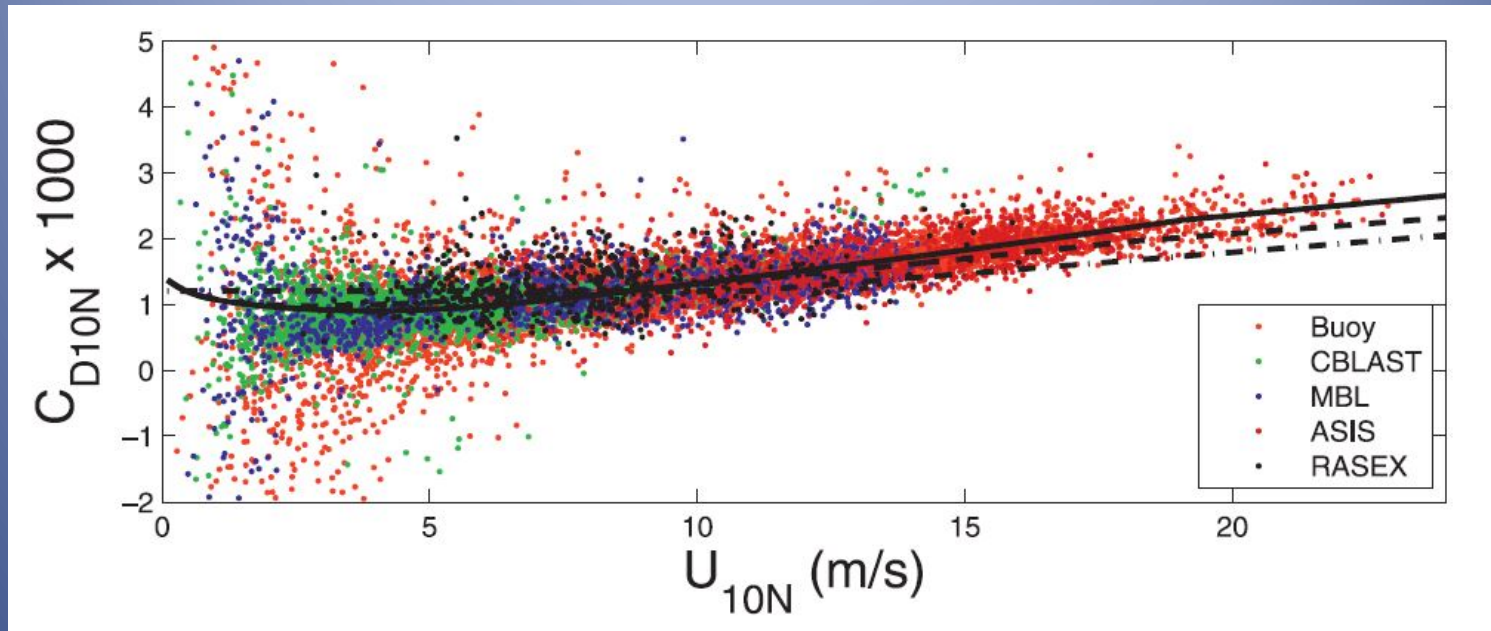
$$S_{\text{ds}}^{\text{loc}}(c) \, \text{dc} = \text{b} \, g^{-2} c^5 \Lambda(c) \, \text{dc}$$

$$\Lambda(c) = S_{\text{ds}}^{\text{loc}}(c) * g^2 / (\text{b} * c^5)$$

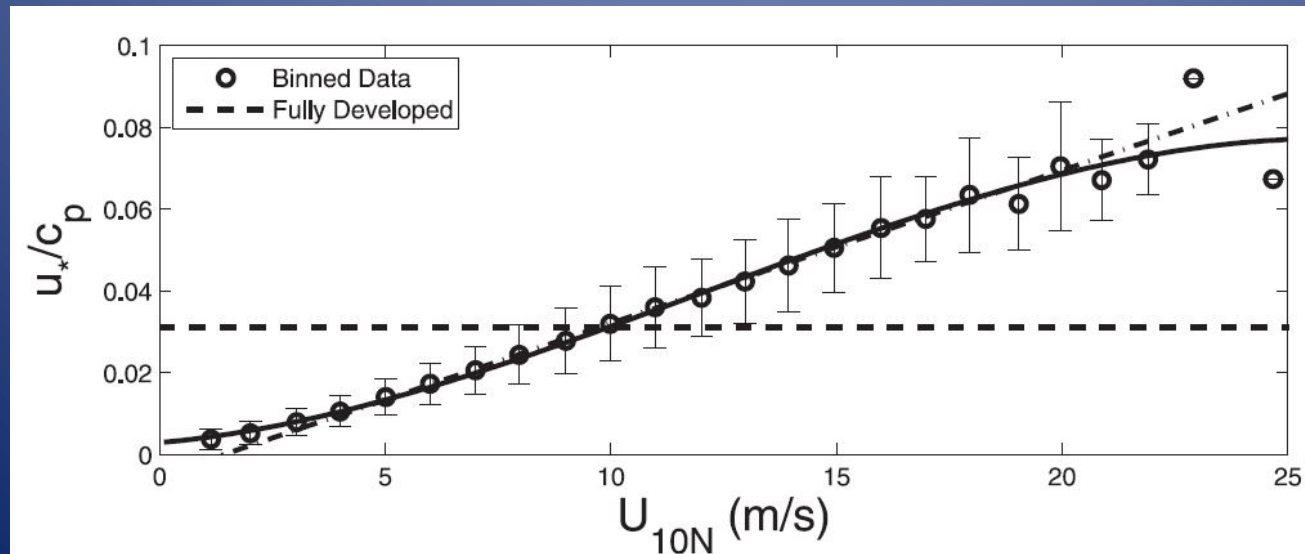
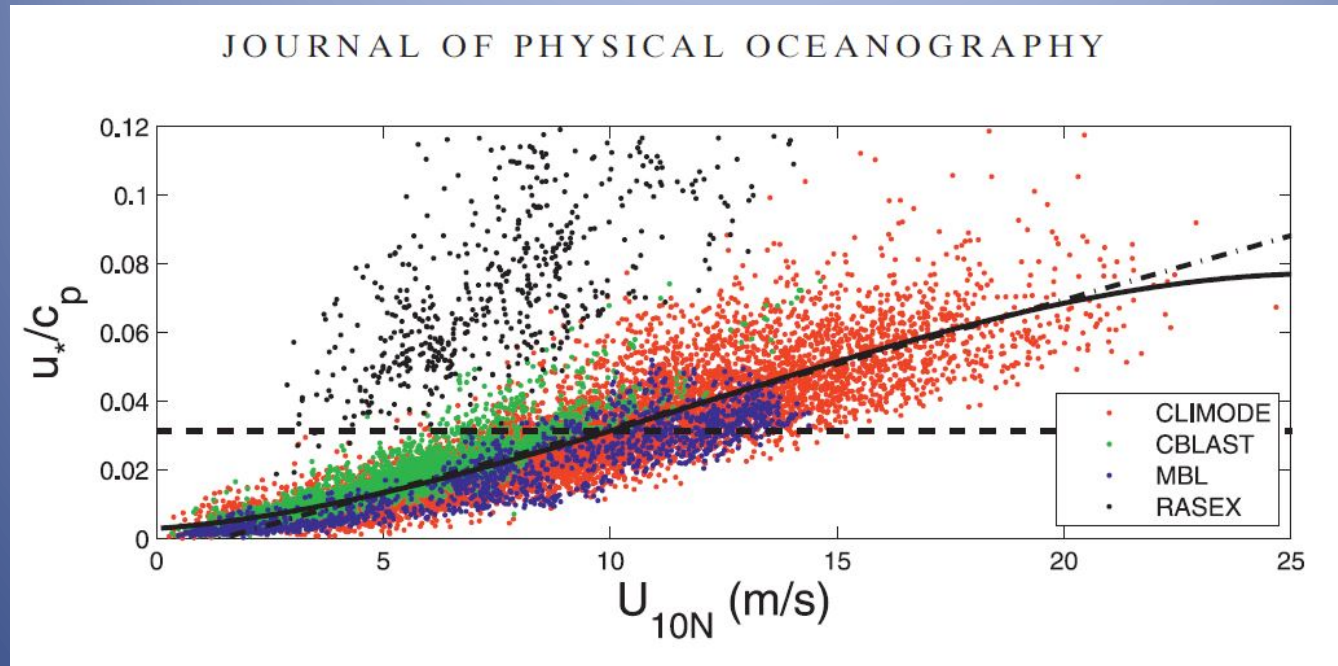
Non-Dimensional Evolution



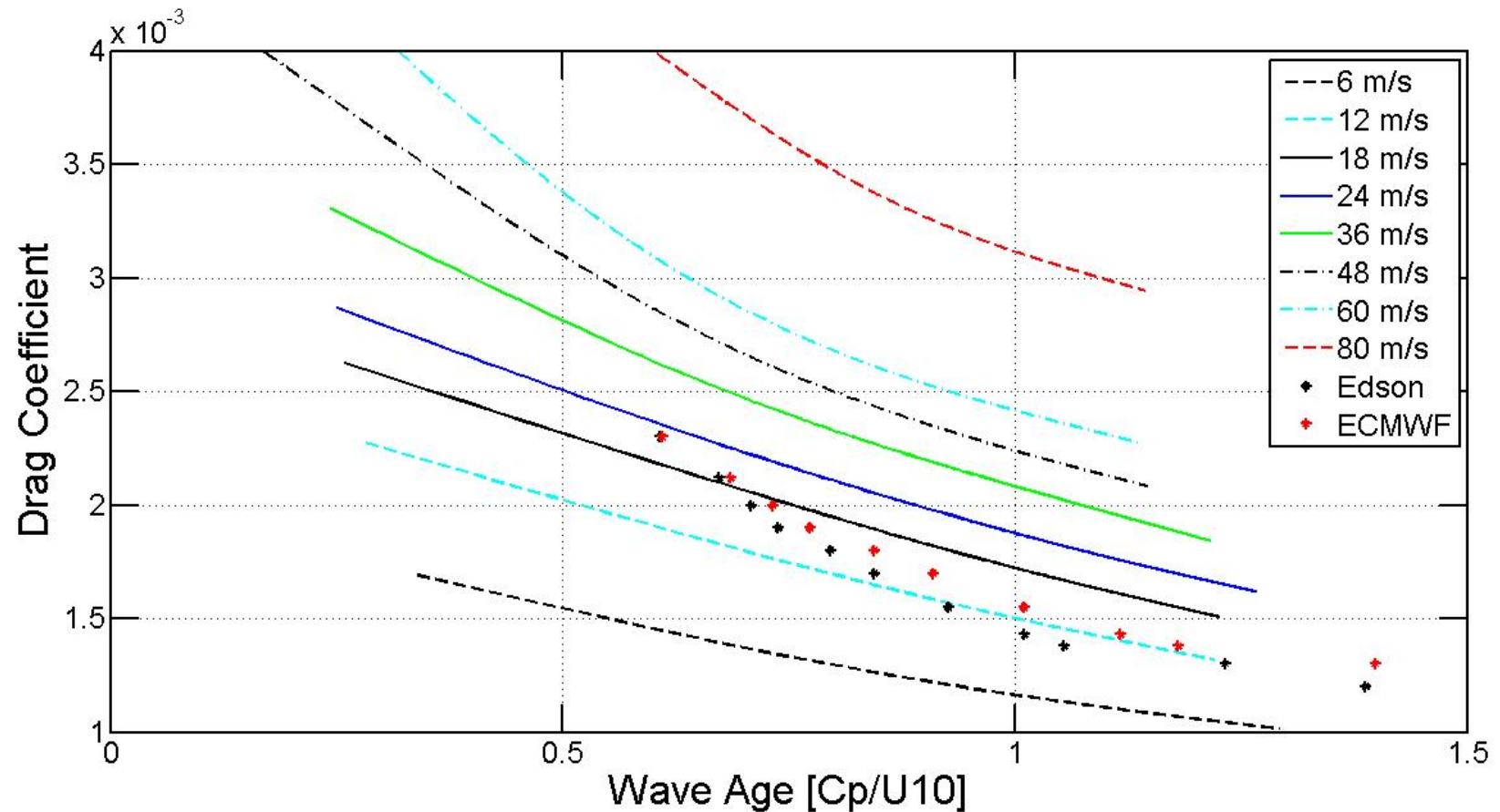
Edson et al, 2013. Drag Coefficient Observational Data.



Edson et al, 2013. Observational Data: Wave Age .vs. U10

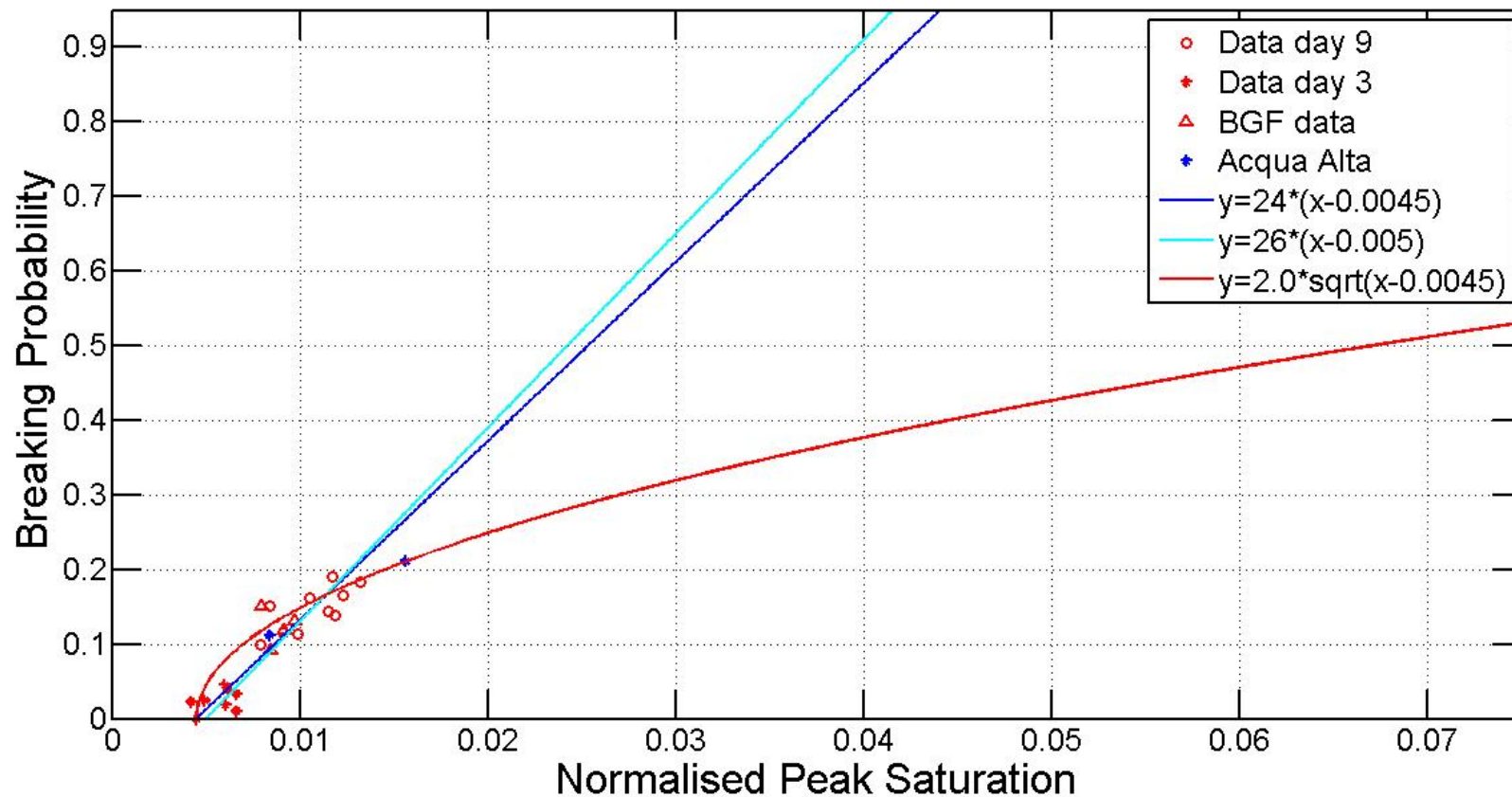


Model Forecast : Drag Coefficient forecast for Fetch Limited for multiple Wind Speeds 6 m/s to 80 m/s

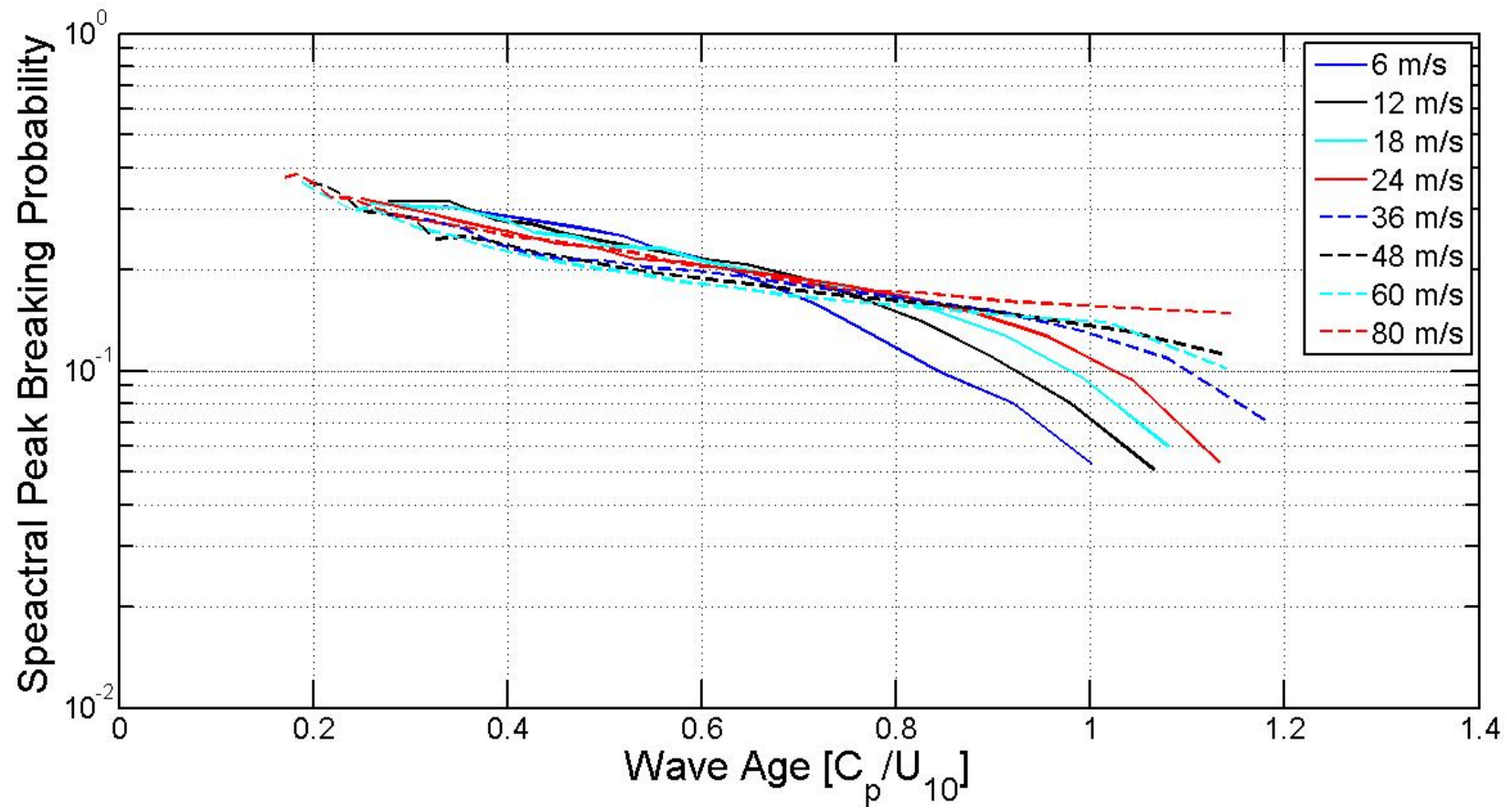


Breaking Probability .vs. Normalized Saturation.

Observed Data; Previous Parameterization; Current Parameterization

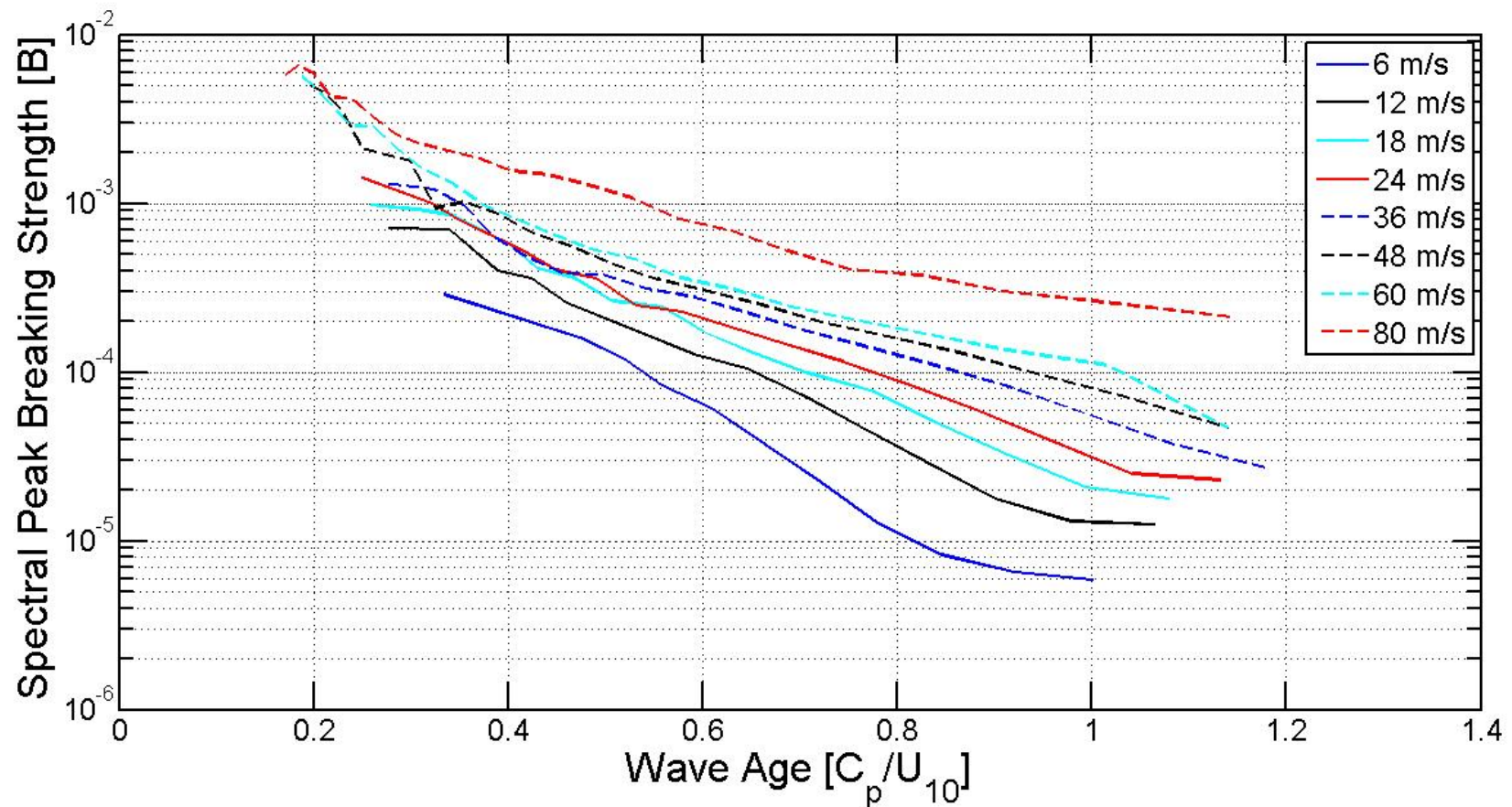


Model Forecast: Breaking Probability .vs. Wave Age for multiple wind speeds 6 m/s to 80 m/s.



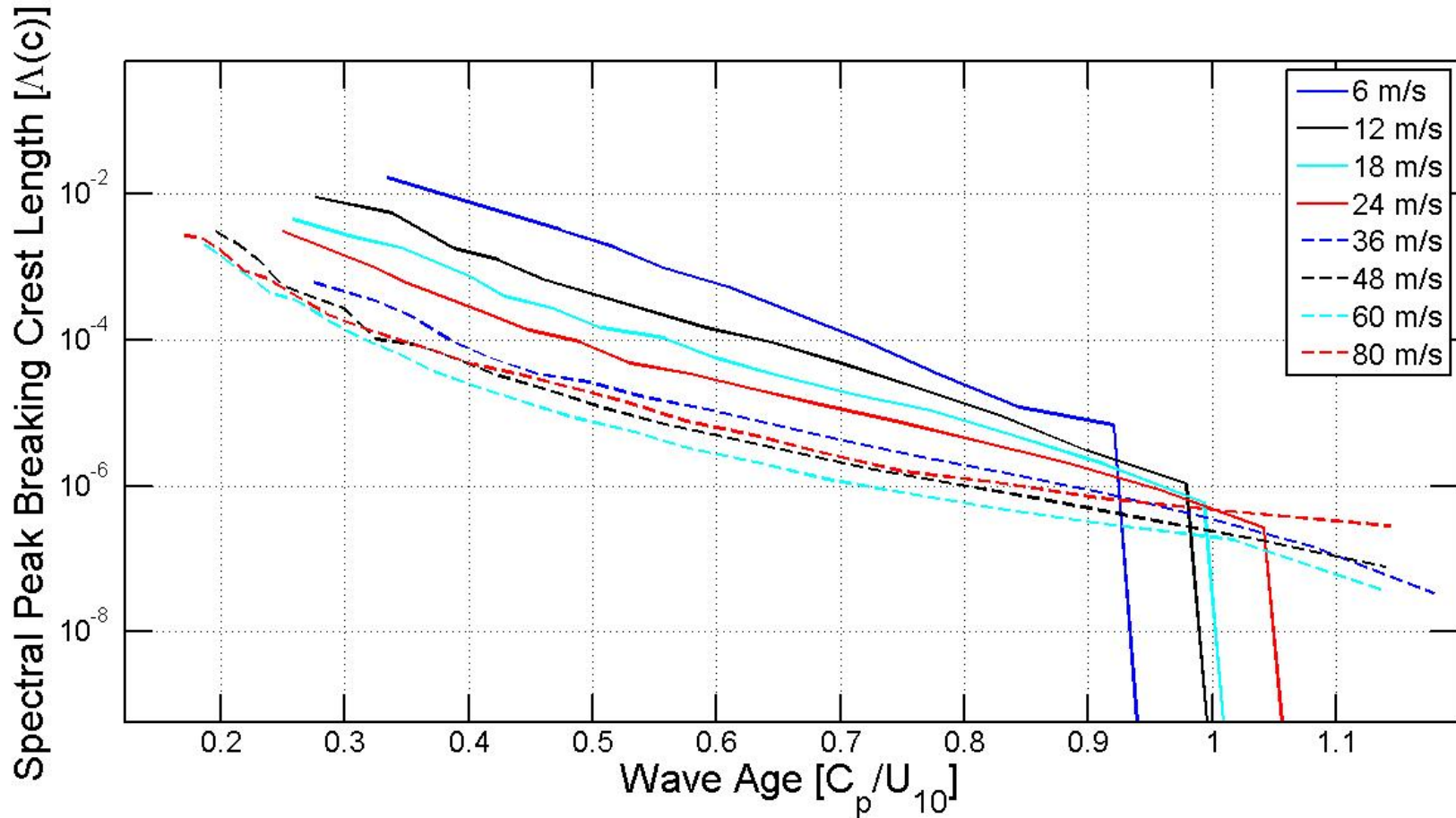
Model Forecast: Breaking Strength .vs. Wave Age

For multiple wind speeds 6 ms/ to 80 m/s.

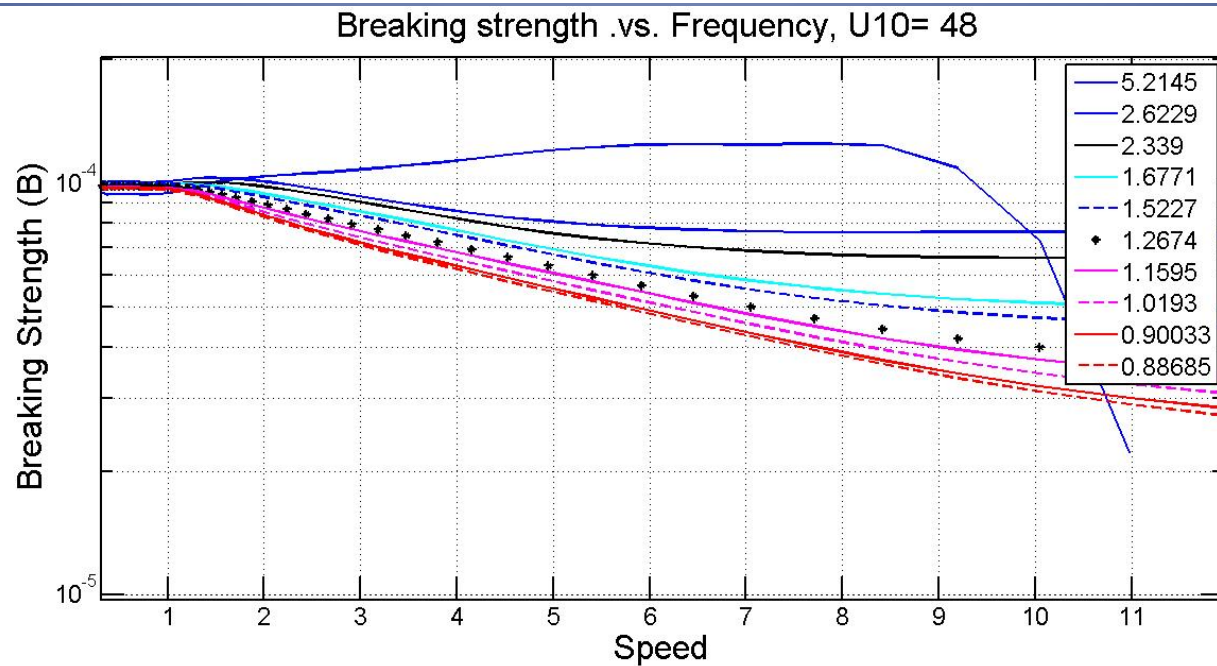
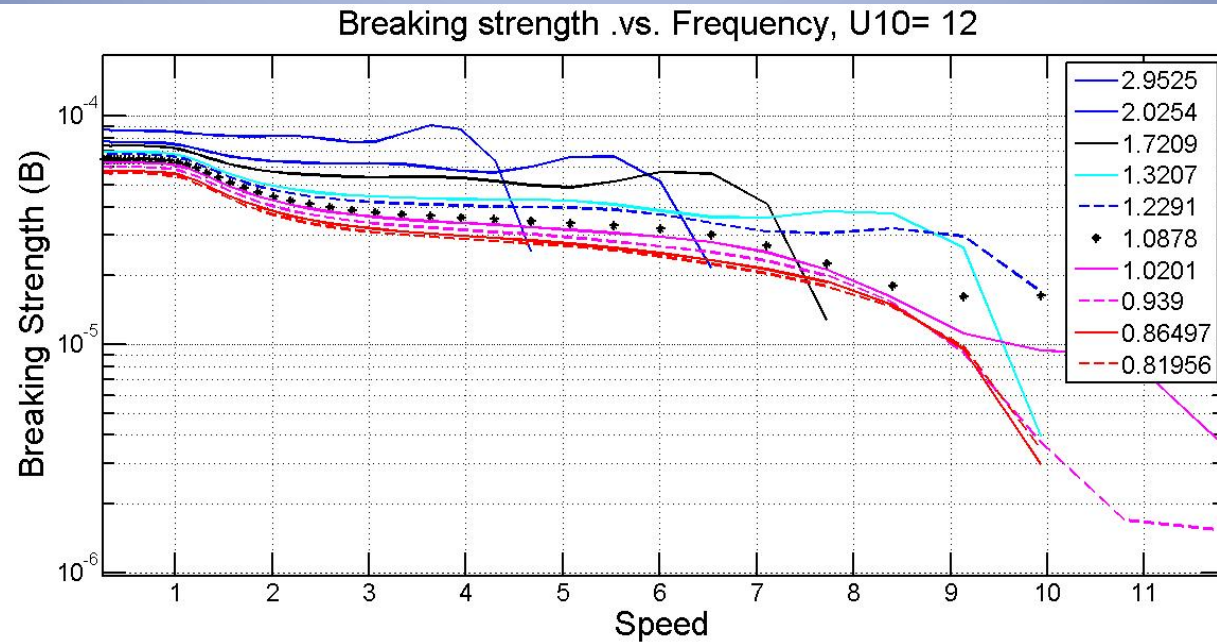


Model Forecast: $\Lambda(c)$.vs. Wave Age.

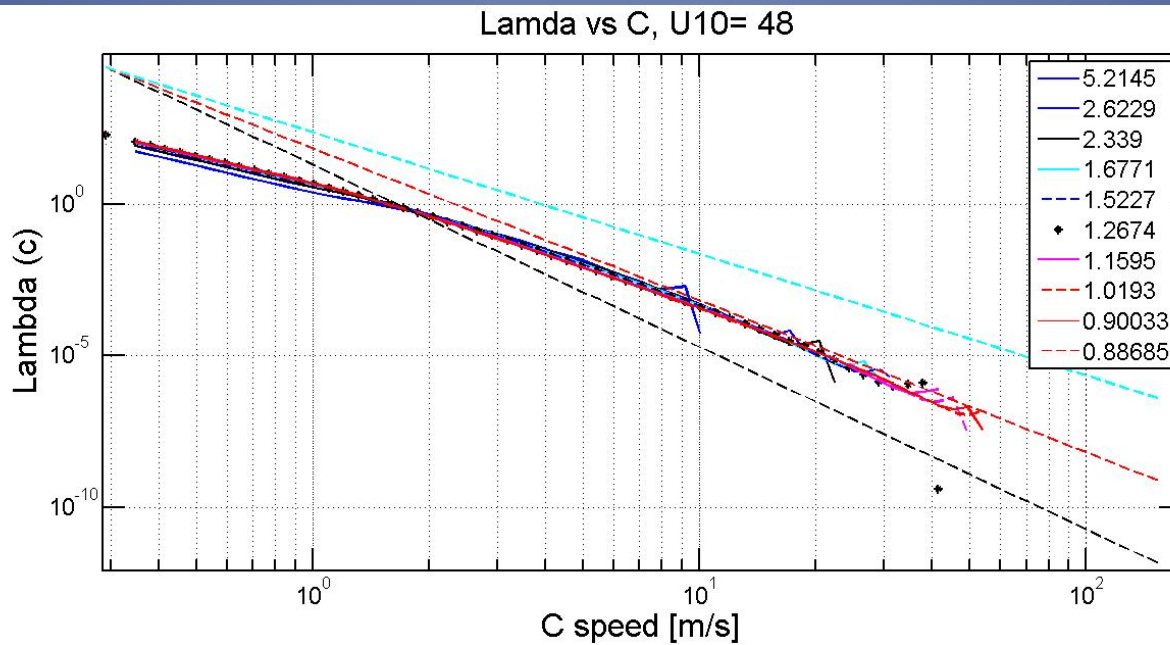
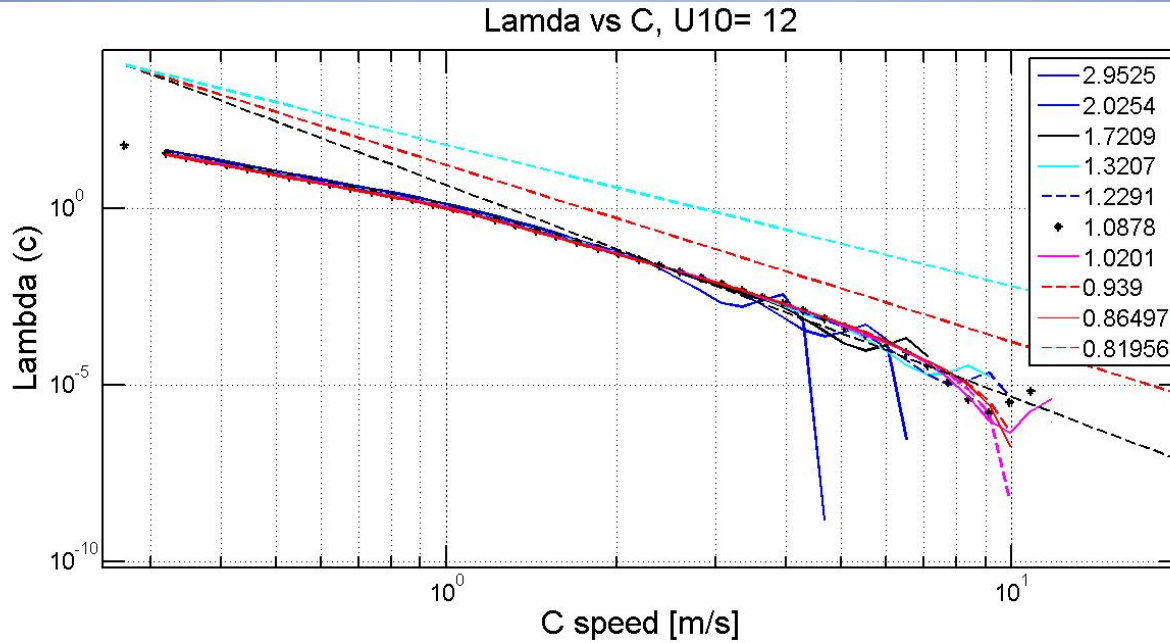
For multiple wind speeds 6 ms/ to 80 m/s.



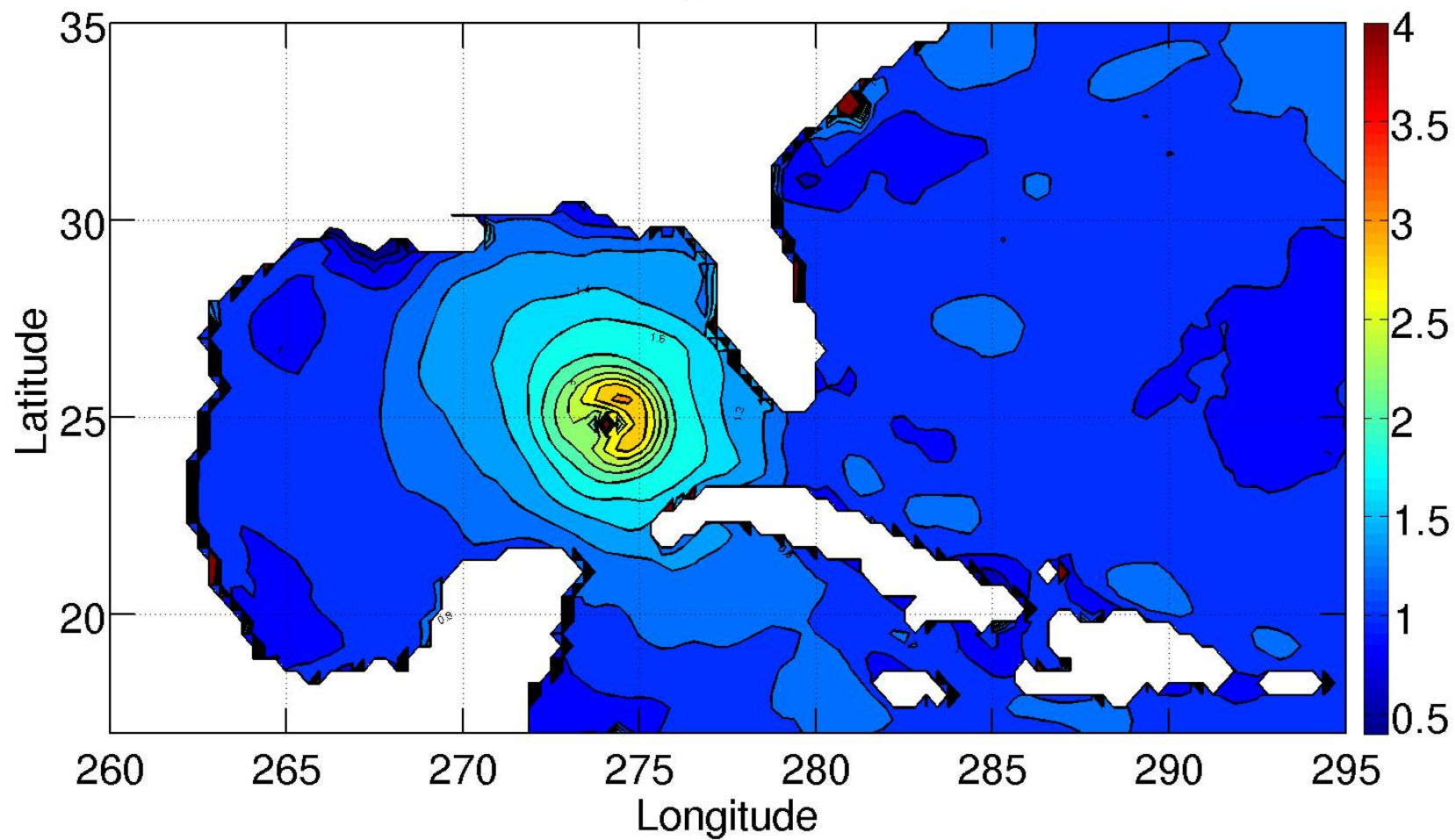
Model Forecast: Breaking Strength [b] .vs. Speed for 12m/s & 48 m/s.



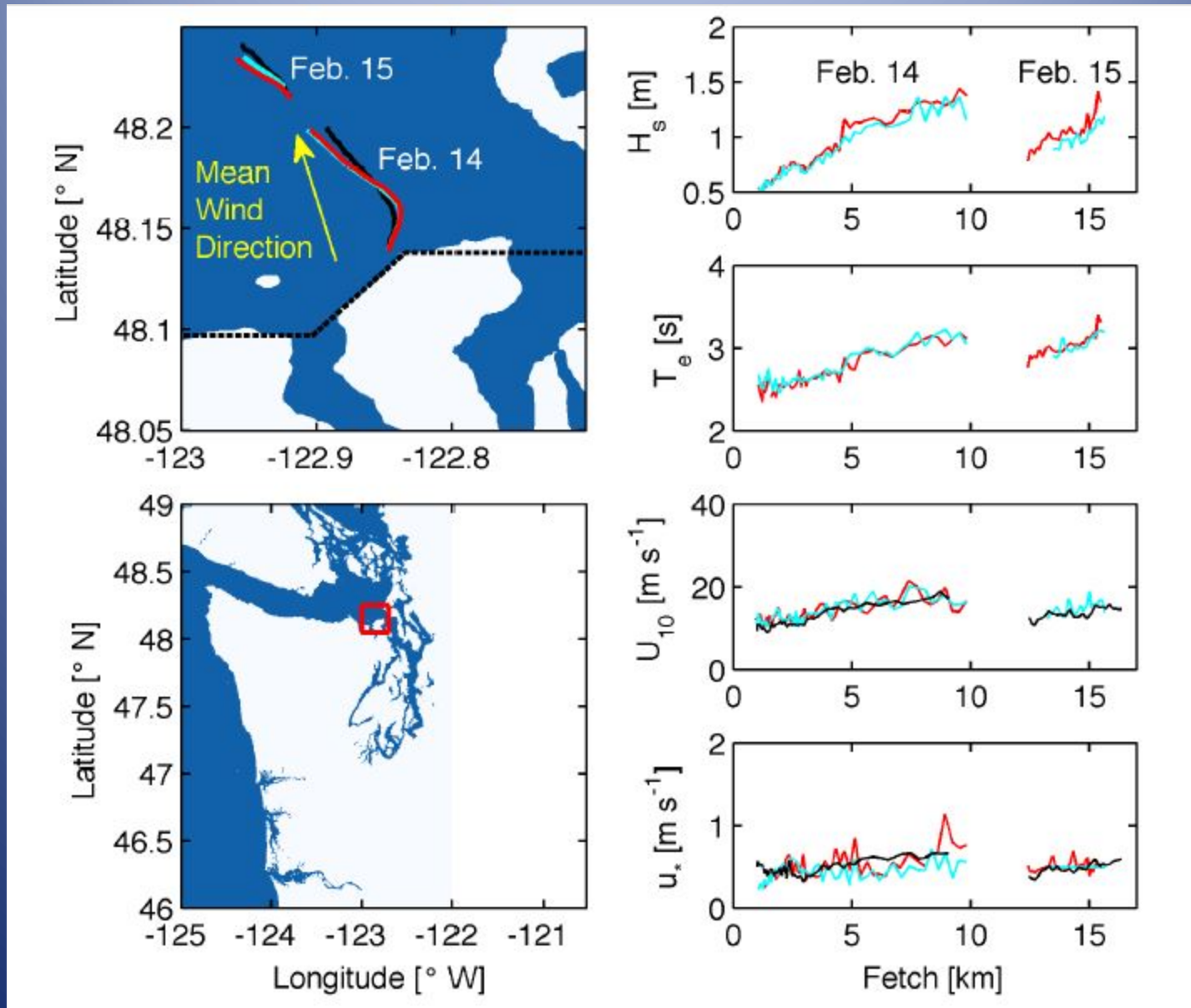
Model Forecast: Lambda(c) .vs. Speed for 12 m/s & 48m/s.



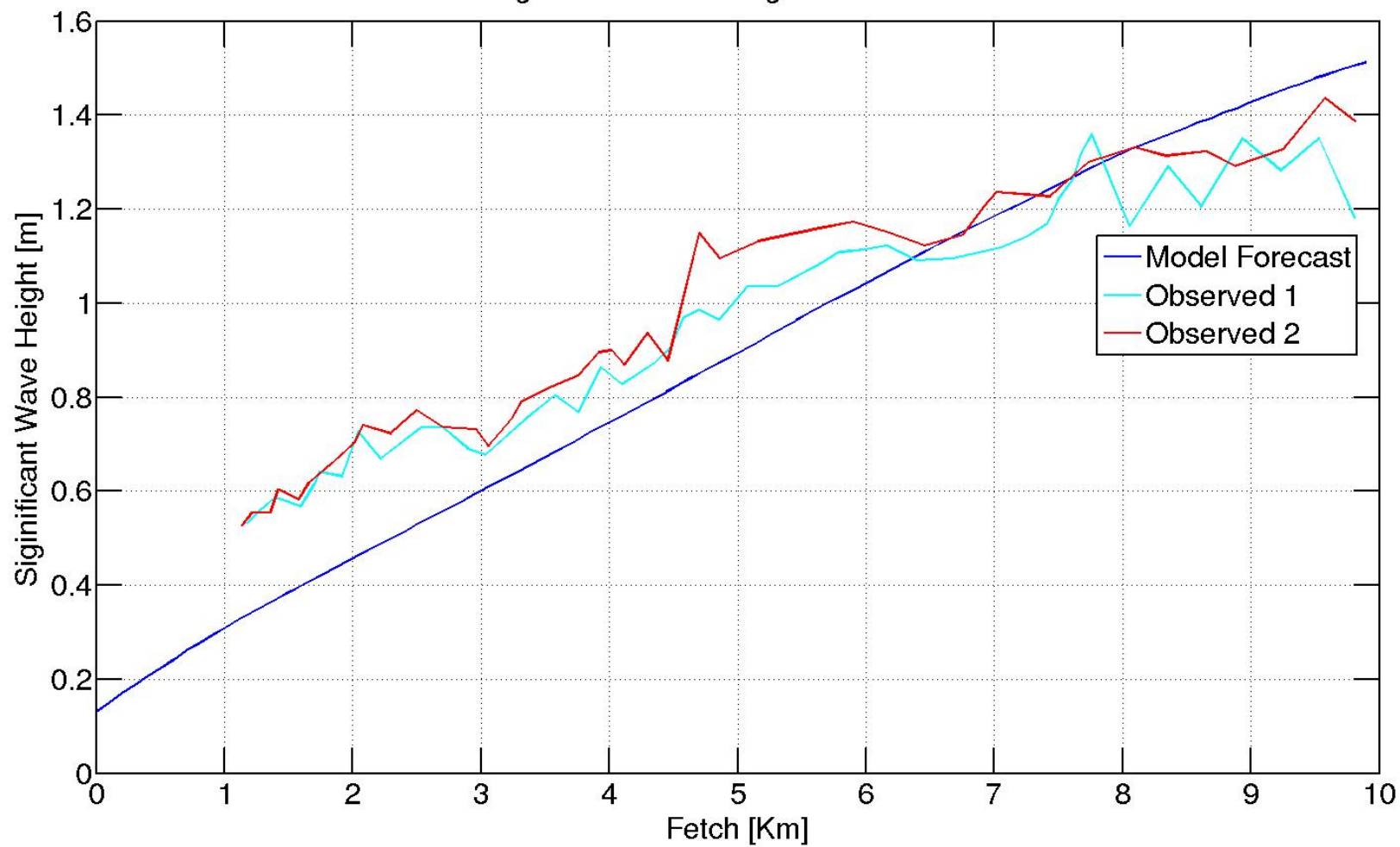
Modeled Drag Coefficient $\times 10^3$

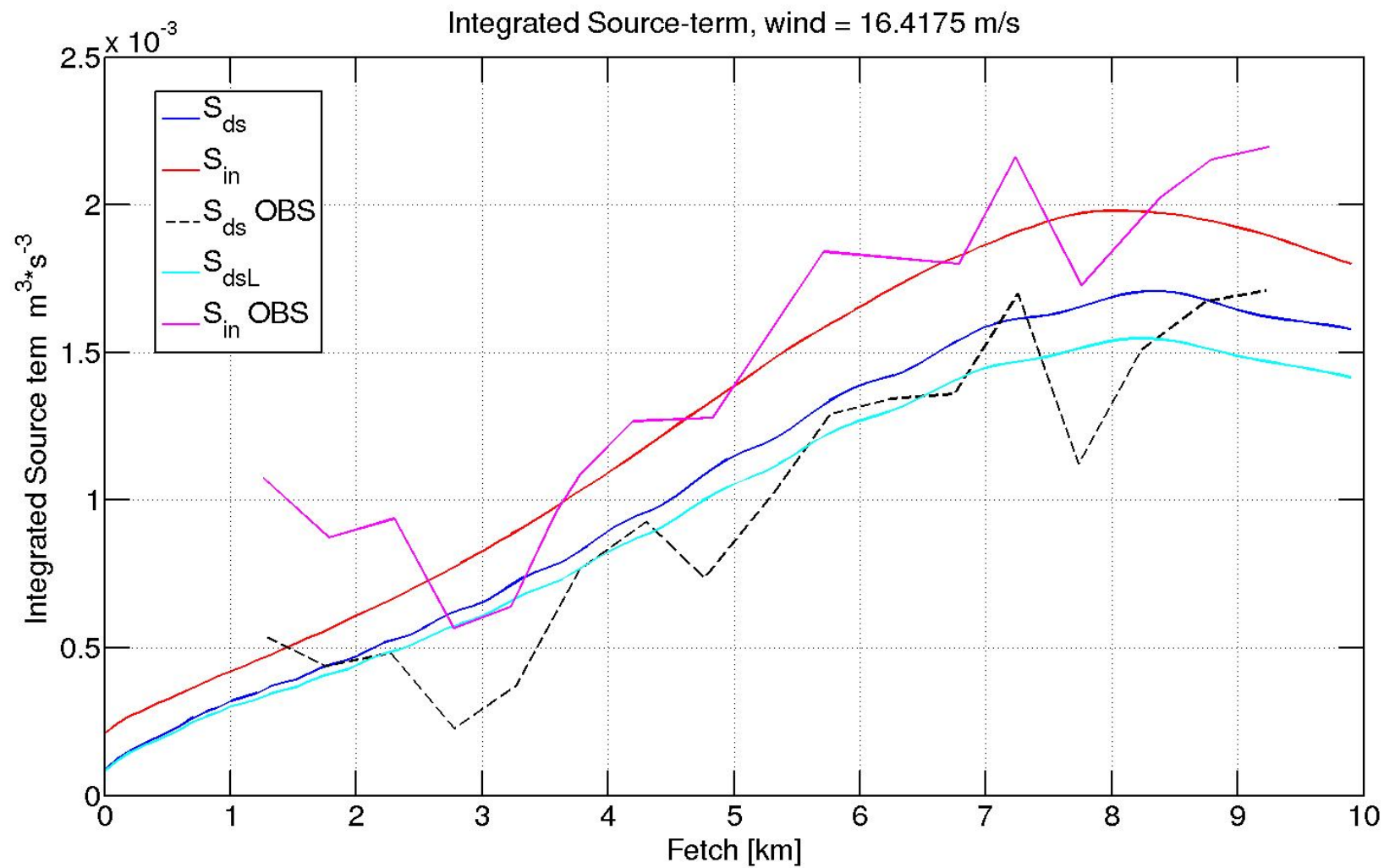


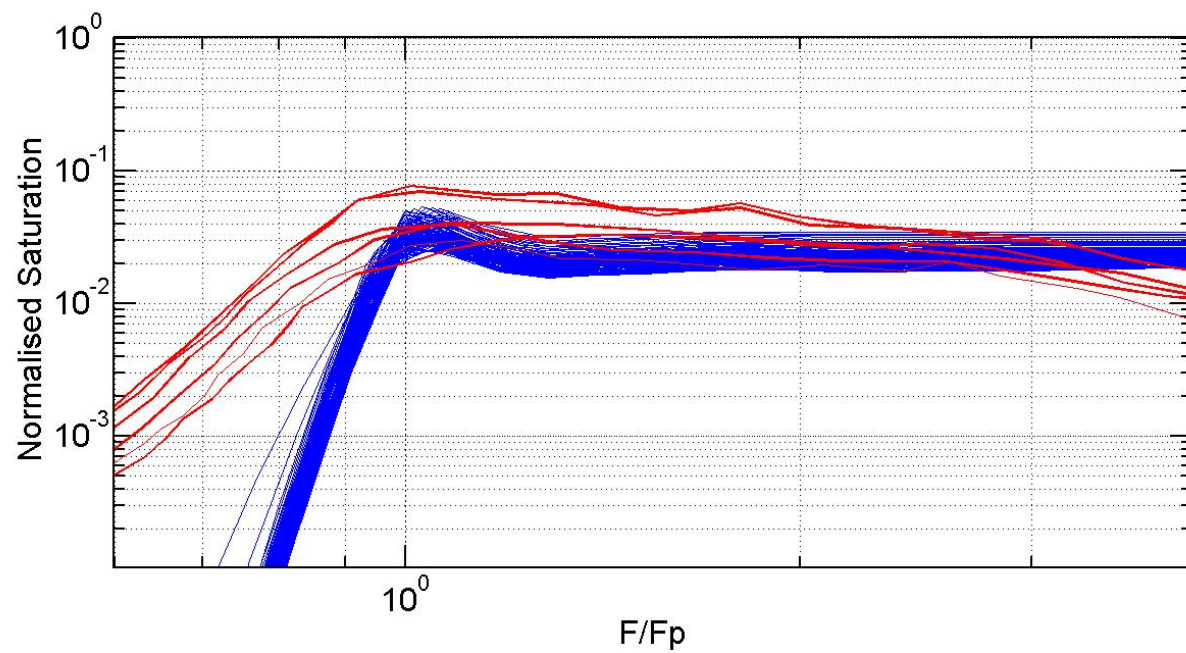
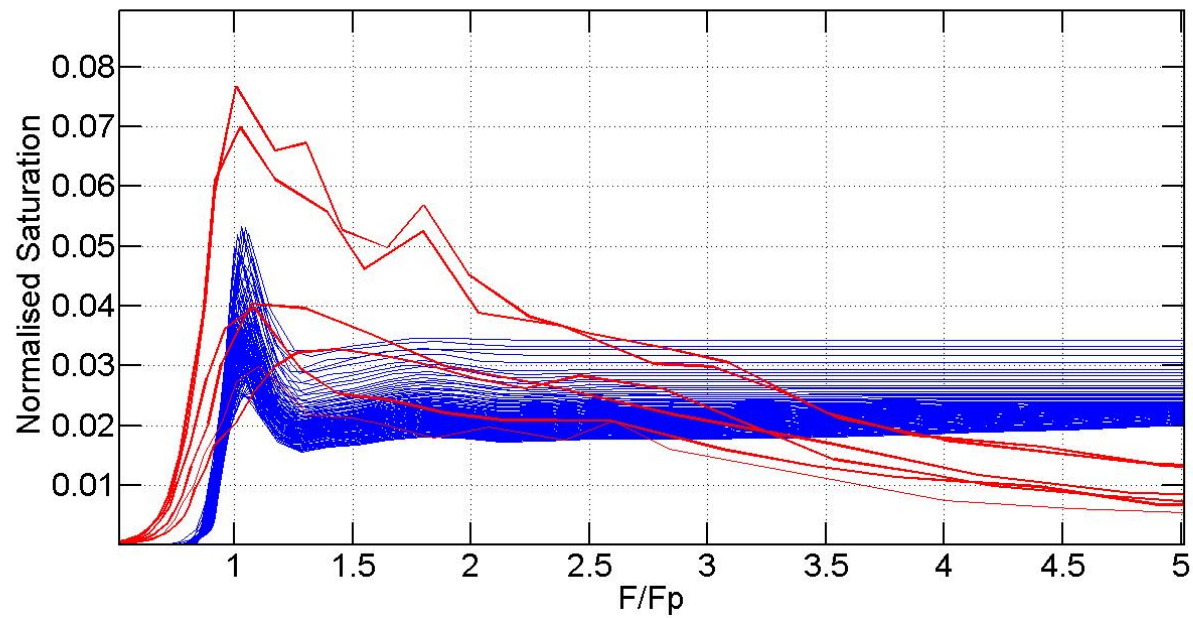
Strait of Juan de Fuca: Experiment Conditions

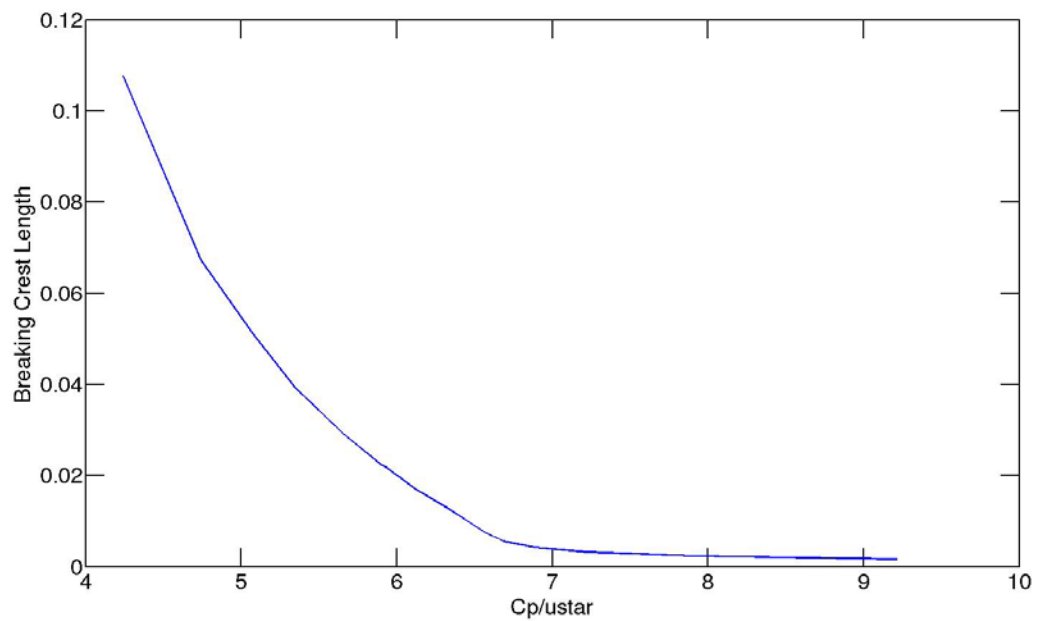
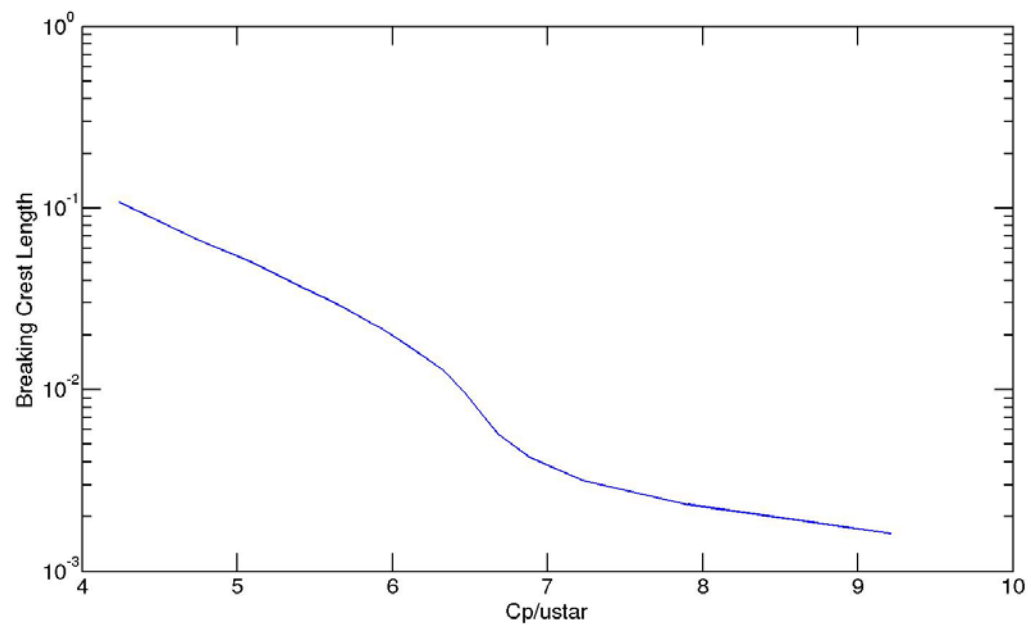


Significant Wave Height versus Fetch









Concluding Remarks

- our framework provides predictions of wave breaking properties (breaking probability, breaking crest length spectral density per unit area and breaking strength) using standard spectral wave models.
- it provides accurate predictions for the limited breaking data available in developing and mature wind seas
- further validation against data will be made as suitable new data sets become available.
- it has been added to existing spectral wave forecasting models. Upgrading the form of the DIA is desirable.

We cant Save Everyone !

