

# Incorporating Breaking Wave Predictions in Spectral Wave Forecast Models

Russel Morison and Michael Banner

School of Mathematics, The University of New South Wales,  
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## Motivational Aspects

- **wave breaking at sea** - widespread air-sea interfacial process with very significant geophysical and maritime importance
- present spectral wave forecast models **do not provide explicit forecasts** of breaking wave properties.
- recent advances in understanding wave breaking have made it possible to redress this deficiency
- to describe a novel methodology that adds to standard spectral wave model output - **accurate forecasts** of
  - (i) the **spectral density of breaking crest length per unit area** and
  - (ii) the associated **breaking strength**

We did this initially for the **dominant wind waves** and have now extended it across the **full spectrum**



## St Jude's Storm shutdown: Eurostar and Monday morning train services cancelled across south as coast is lashed by 25ft waves

- Amendments and cancellations on First Capital Connect, Southeastern, Greater Anglia and Stansted Express
- Also disruption on East Coast, c2c, First Great Western, Southern, Gatwick Express and South West Trains
- Ferries from Poole and Weymouth to Guernsey & Jersey cancelled and hovercrafts to Isle of Wight suspended
- About 60 flights cancelled at London Heathrow Airport tomorrow but none yet at Gatwick, Stansted and Luton
- Forecasters warn houses face damage, trees falling and power cuts in biggest storm to hit Britain in a decade
- Wales and South West England will be hit first early tomorrow morning with winds of up to 90mph expected
- Boy, 14, believed to have drowned today after swimming with friends in waves off Newhaven in East Sussex
- Canoeist dies after being pulled from swollen River Tees near Barnard Castle, County Durham, after capsizing

By [MARK DUELL](#), [TARA BRADY](#) and [JONATHAN PETRE](#)





# Modeling Wave Breaking Spectrally

Radiative transfer equation (deep water, no currents)

The radiative transfer equation for describing the evolution of the wave height spectrum  $F(k)$  is given by:

$$\frac{\partial F}{\partial t} + c_g \cdot \nabla F = S_{in} + S_{nl} + S_{diss}$$

where

- $F=F(k, \theta)$  is the directional wave spectrum
- $c_g$  is the group velocity
- $S_{tot} = S_{in} + S_{nl} + S_{ds}$  is the total source term.
- $S_{in}$  is the **atmospheric input** spectral source term
- $S_{nl}$  is the **nonlinear spectral transfer** source term representing nonlinear wave-wave interactions within the spectrum
- $S_{ds}$  is the spectral **dissipation rate** here taken as due primarily to wave breaking

# Saturation Threshold-based Dissipation Rate $S_{ds}$

- based on treating spectral bands as nonlinear wave groups. Uses a low power of the spectral saturation ratio (~steepness ratio) to simulate observed threshold behaviour [*extension of Alves & Banner (JPO, 2003)*]

$$S_{ds}(k, \theta) = [C_1 * D * (\tilde{\sigma} - \tilde{\sigma}_T) / \tilde{\sigma}_T]^{a_1} + C_2 * D * E_{tot} * k_p^2] (\sigma / \sigma_m)^{a_2} \omega F(k, \theta)$$



‘local  $S_{ds}$ ’



‘non-local  $S_{ds}$ ’

## This formulation uses

- normalized azimuthally-integrated saturation:  $k^4 F(k) / \theta(k) = (2\pi)^4 f^5 F(f) / 2g^2$
- measured threshold of the normalized spectral saturation (Banner et al., JPO, 2002) with  $a_1=2$
- tail exponent  $a_2 = 4$  to match dissipation to input behavior in the spectral tail
- nonlocal dissipation rate component
- coefficient D for the local  $S_{ds}$ : non-dimensional and linear in the wind speed to match to the wind input term.
- $C_1$  and  $C_2$  constants

# Modified Jansen Wind Input

$$S_{in}(k, \theta) = \varepsilon \beta(k, \theta) \omega (u_*^{red}(k) \cos \theta / c)^2 * E(k, \theta)$$

$$\beta(k, \theta) = J_2 \mu (\ln(\mu))^4 / \kappa^2 \quad \text{where } J_2 = 1.6 \quad (\text{Janssen (1991) used } 1.2).$$

$$\beta(k, \theta) = 0 \quad \text{for} \quad \mu > 1$$

$$\mu(k, \theta) = (u_*/c)^2 (g z_0 / u_*^2) \exp(J_1 \kappa / (u_* \cos \theta / c)^2)$$

$$z_0 = \frac{0.01 u_*^2}{g} / \sqrt{1 - C_0(\tau_w / \tau)}$$

$$u_*^{red}(k_n) = \sqrt{[\tau_{tot} - J_0 \sum_{i=1}^n (\tau_w(i) + \tau_{bw}(i))] / \rho_{air}}$$

# Brief description of the methodology

$\Lambda(c)$  is the spectral density of breaking wave crest length per unit area

$\Pi(c)$  is the spectral density of the total wave crest length per unit area

The breaking probability  $P_{br}(c)$  for wave scales  $c$  is defined as:

$$Pr_{br}(c) = \frac{\int c \Lambda \, dc}{\int c \Pi \, dc}$$

~

passage rate of breaking wave crests  
—————  
passage rate of wave crests

$\Lambda(c)$ : spectral density of *breaking wave crest length* per unit area with velocities in the range  $(c, c+dc)$  (Phillips, 1985)

$$b \frac{\rho}{g} c^5 \Lambda(c) dc$$

wave energy dissipation rate at scale  $c$

$$b \frac{\rho}{g} c^4 \Lambda(c) dc$$

momentum flux from waves of scale  $c$  to currents

The sea state threshold variable used for breaking probability was the normalised spectral saturation

$$\tilde{\sigma}(k) = \sigma(k) / \langle \theta(k) \rangle$$

where  $\sigma(k)$  is the azimuth-integrated spectral saturation given by

$$\sigma(k) = k^4 \Phi(k)$$

$$= (2\pi)^4 f^5 G(f) / 2g^2$$

and  $\langle \theta(k) \rangle$  is the mean spectral spreading width given by

$$\langle \theta(k) \rangle = \int_{-\pi}^{\pi} (\theta - \bar{\theta}) F(k, \theta) k d\theta / \int_{-\pi}^{\pi} F(k, \theta) k d\theta$$

where  $\bar{\theta}$  is the mean wave direction, and  $\Phi(k)$ ,  $G(f)$  and  $F(k, \theta)$  are, respectively, the spectra of wave height as a function of scalar wavenumber, frequency and vector wavenumber.

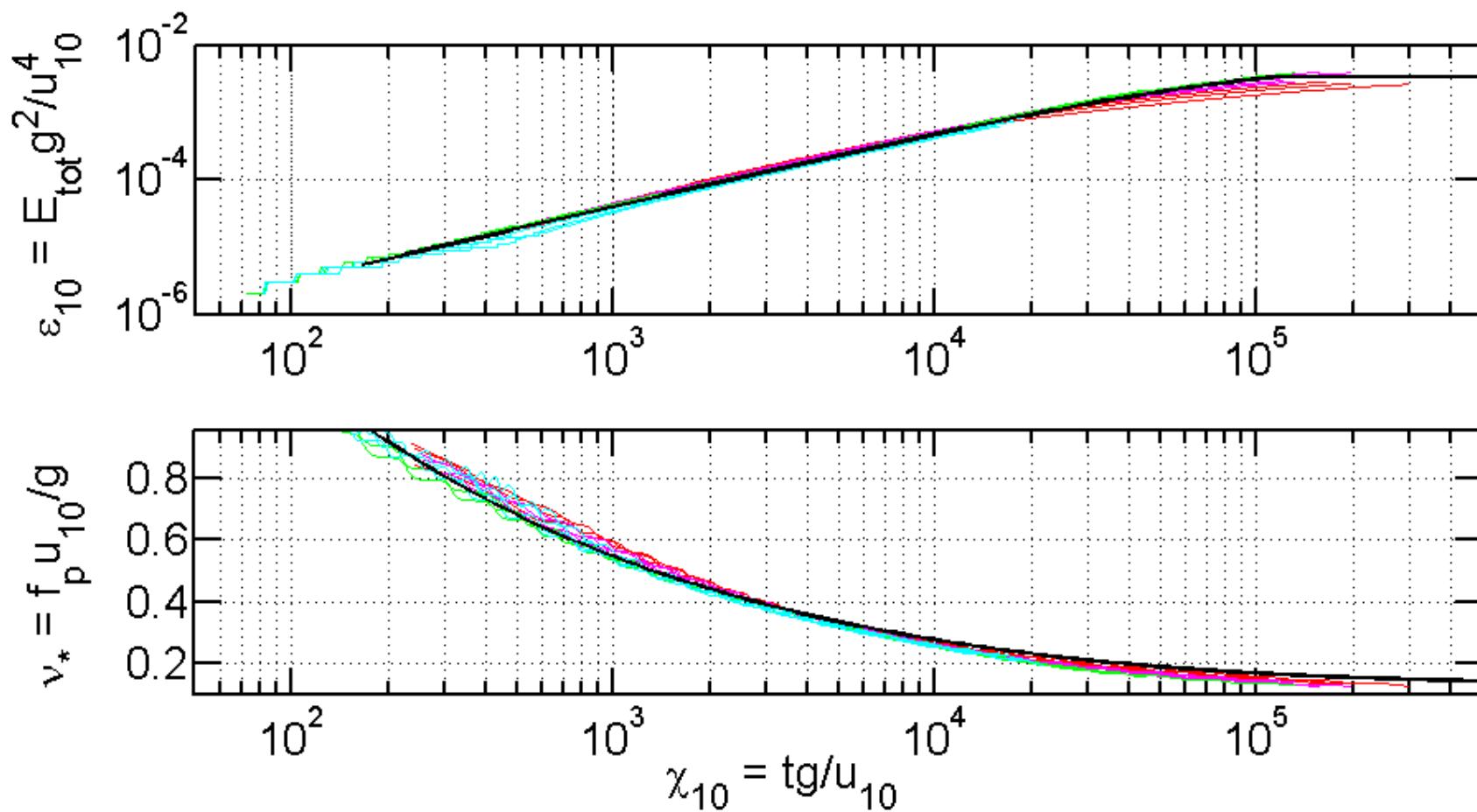
$$\Pr_{\text{br}}(\widetilde{\sigma}) = H(\widetilde{\sigma} - \widetilde{\sigma}_T) * \alpha_{\text{br}} * \big(\widetilde{\sigma} - \widetilde{\sigma}_T\big)^{0.5}$$

$$b_{\text{br}}(\widetilde{\sigma}) = H(\widetilde{\sigma} - \widetilde{\sigma}_T) * c_{\text{br}} * \big(\widetilde{\sigma} - \widetilde{\sigma}_T\big)^{1.0}$$

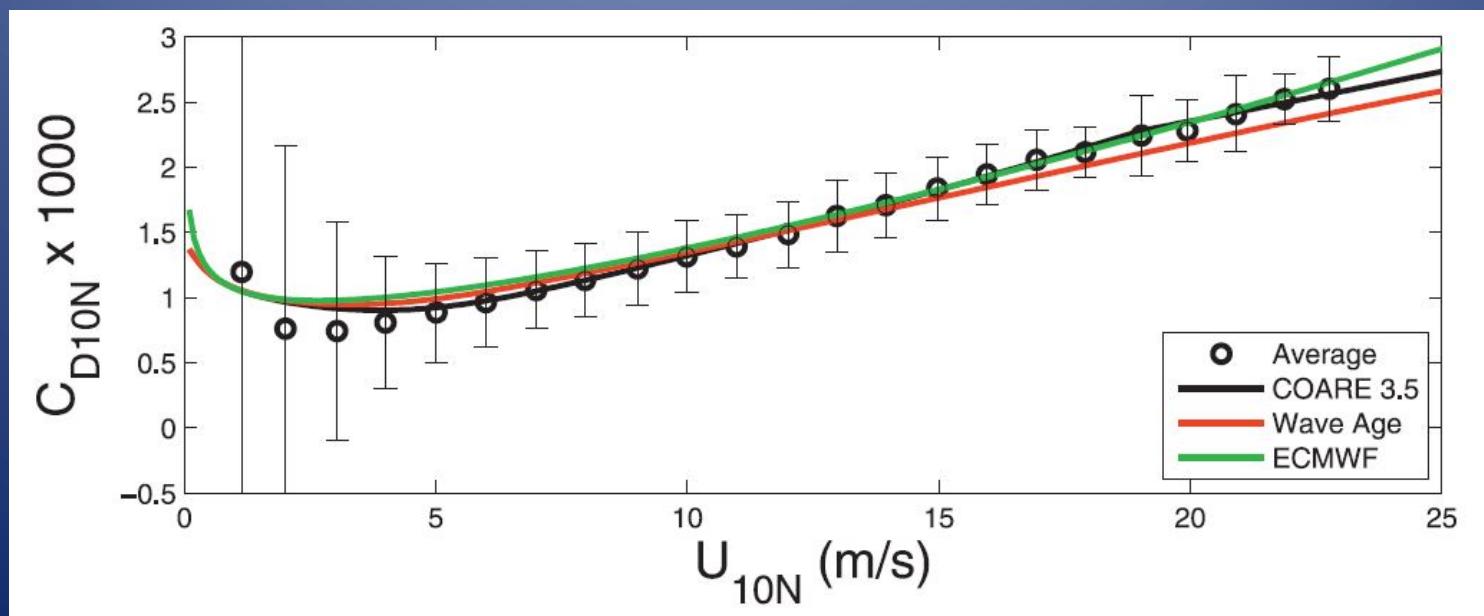
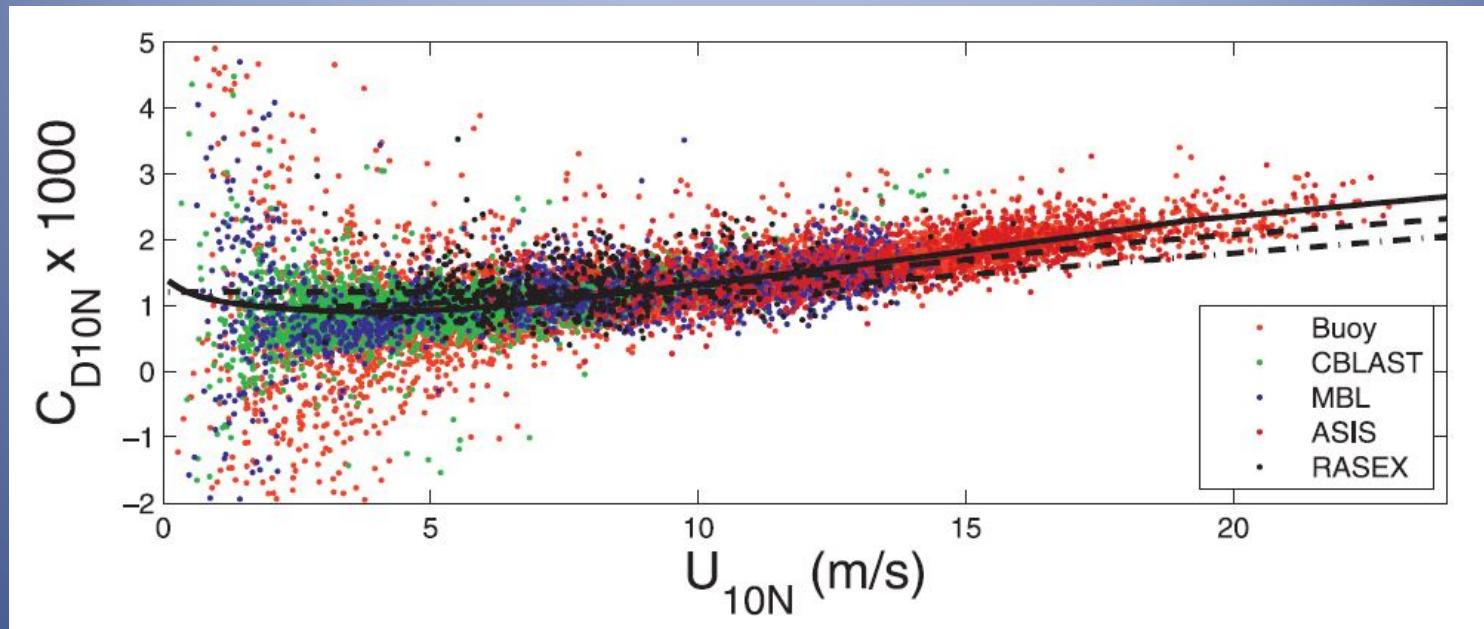
$$S_{ds}^{\rm loc}(c)\,dc = b\;g^{-2}c^5\Lambda(c)\,dc$$

$$\Lambda(c) = S_{ds}^{\rm loc}(c)*g^2/(b*c^5)$$

# Non-Dimensional Evolution

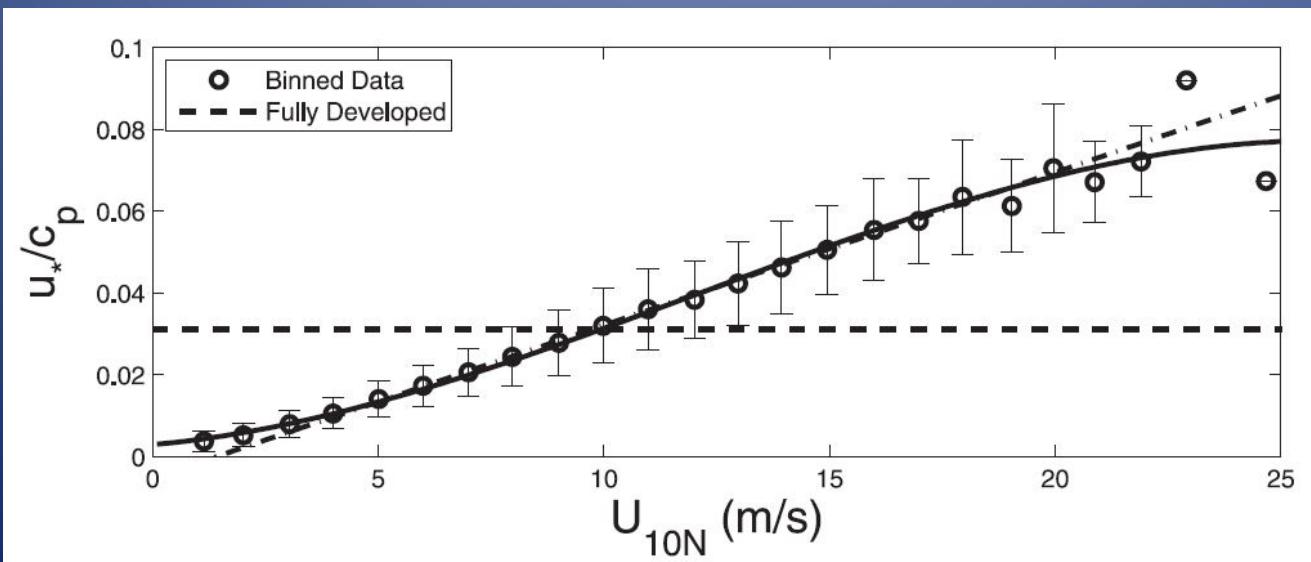
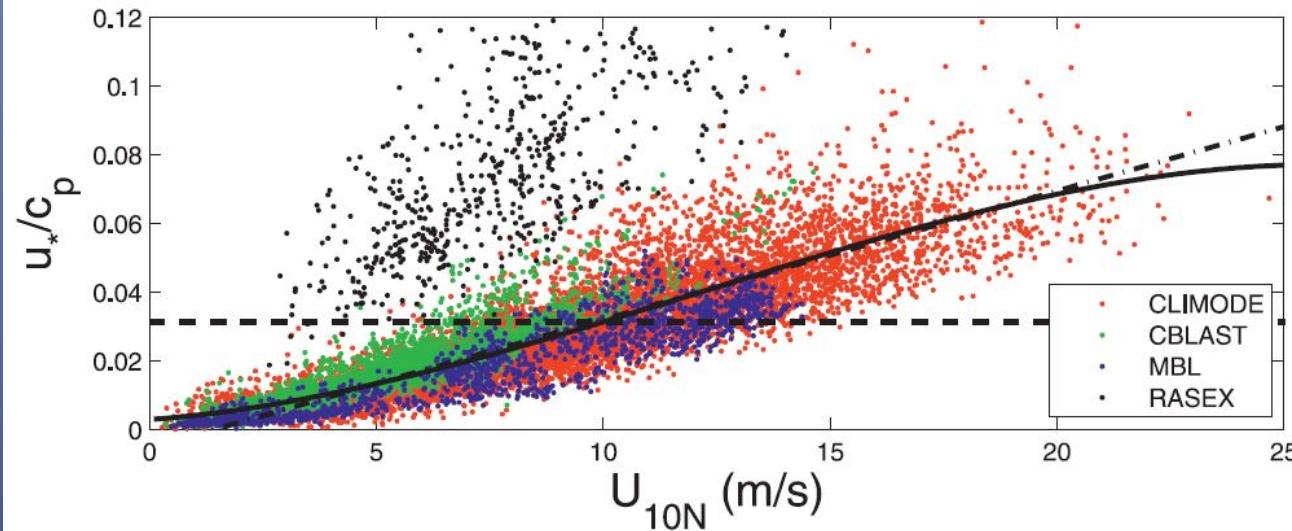


# Edson et al, 2013. Drag Coefficient Observational Data.

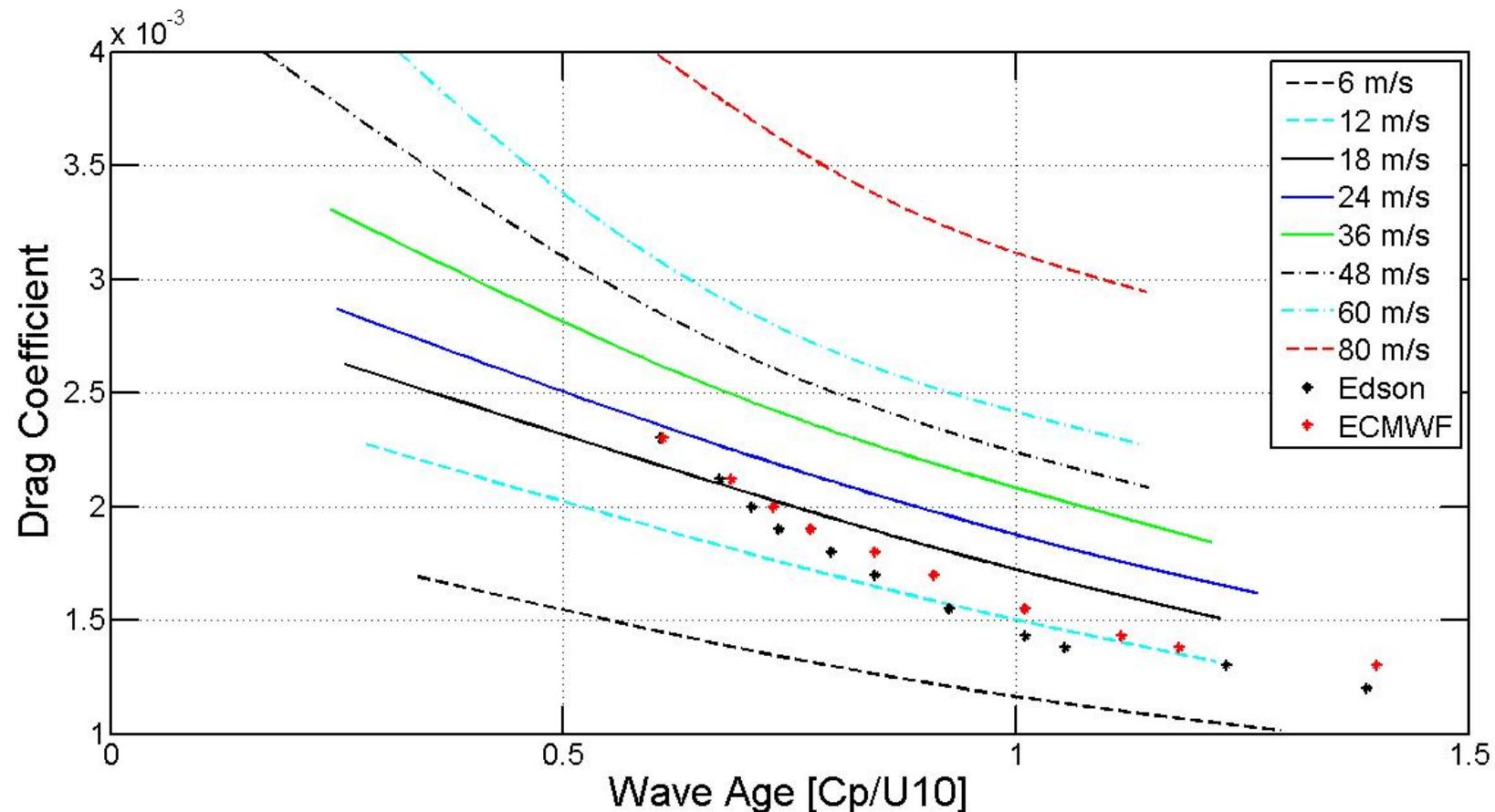


# Edson et al, 2013. Observational Data: Wave Age .vs. U10

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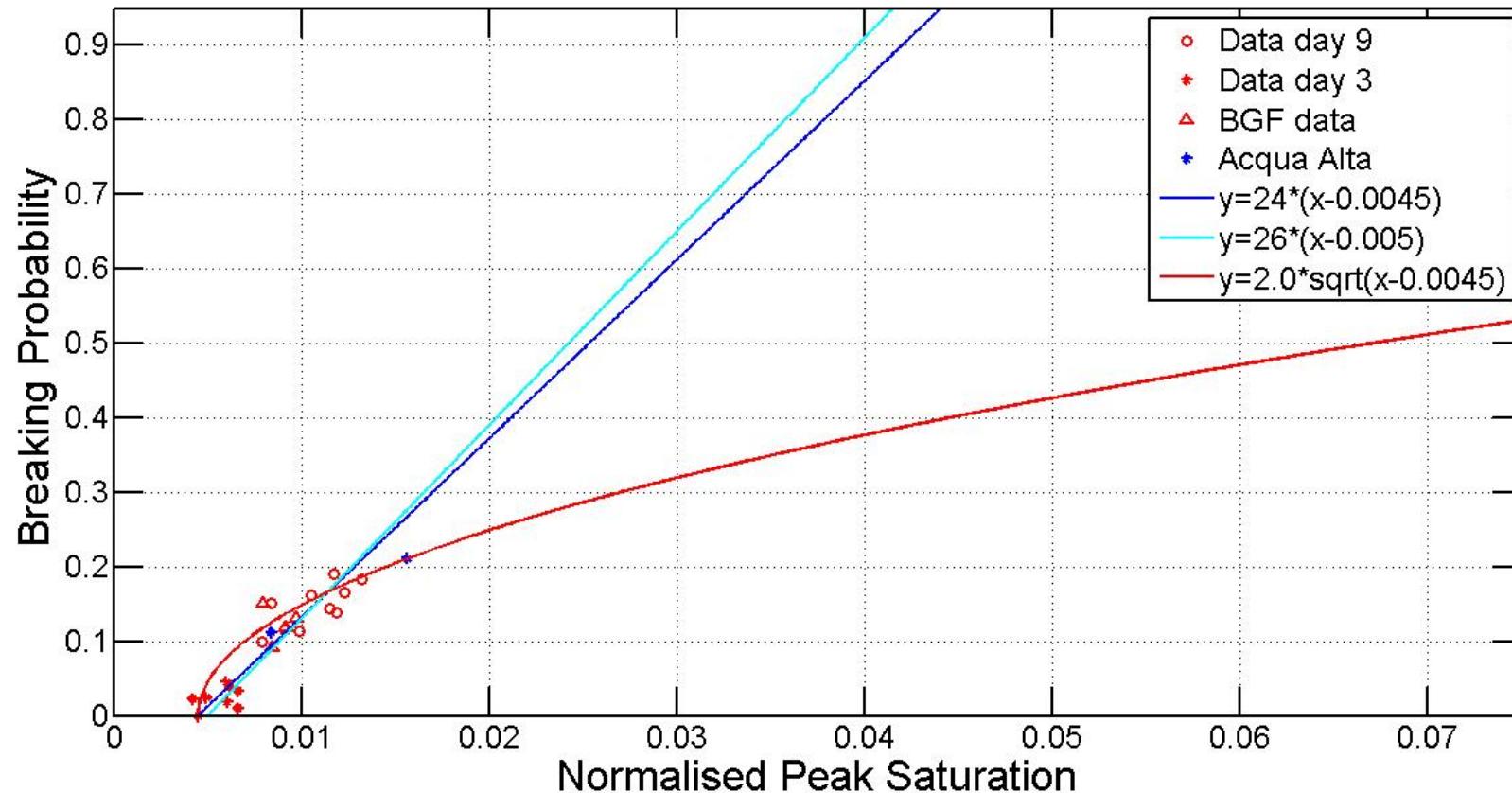


# Model Forecast : Drag Coefficient forecast for Fetch Limited for multiple Wind Speeds 6 m/s to 80 m/s

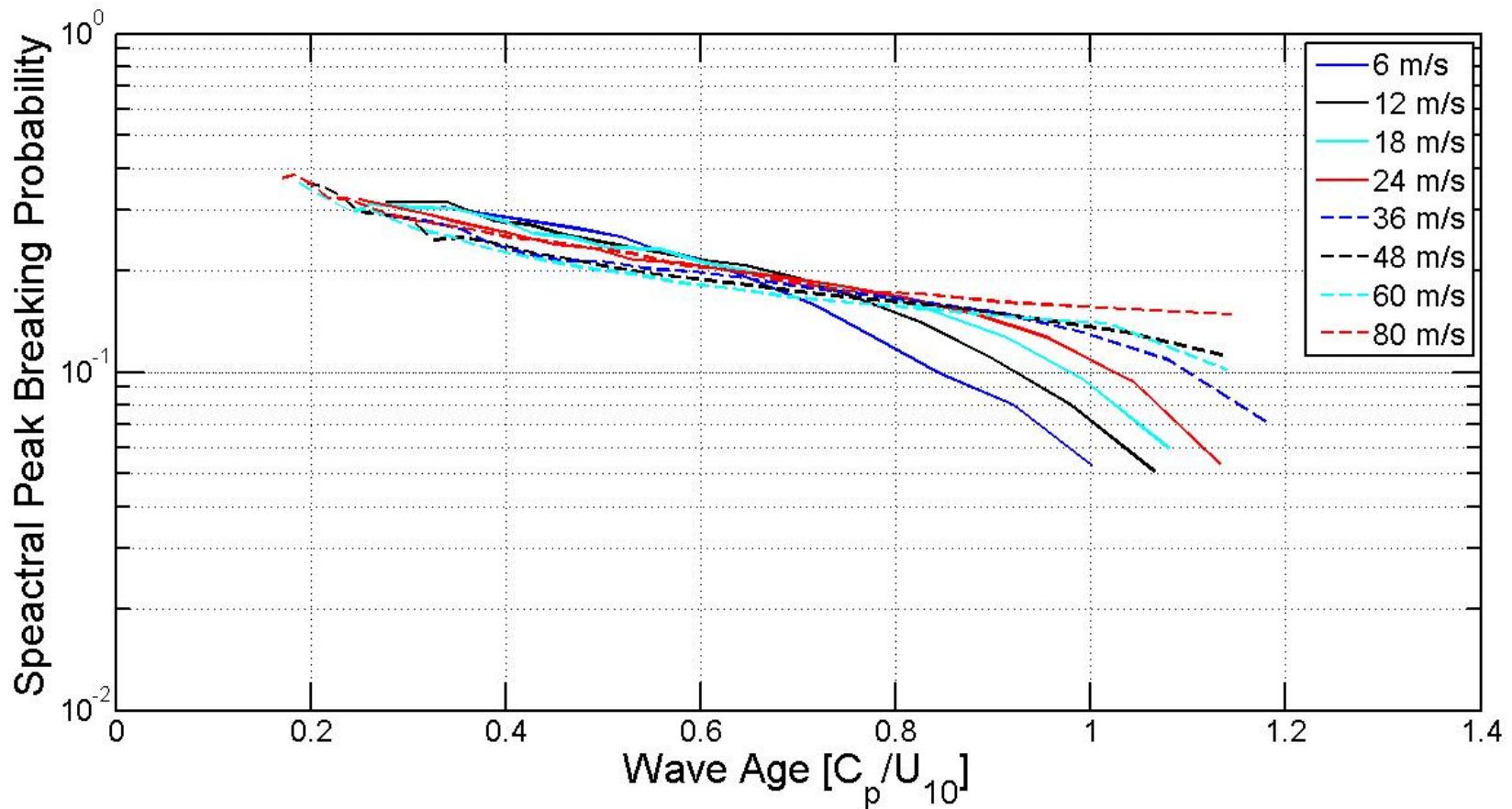


# Breaking Probability .vs. Normalized Saturation.

## Observed Data; Previous Parameterization; Current Parameterization

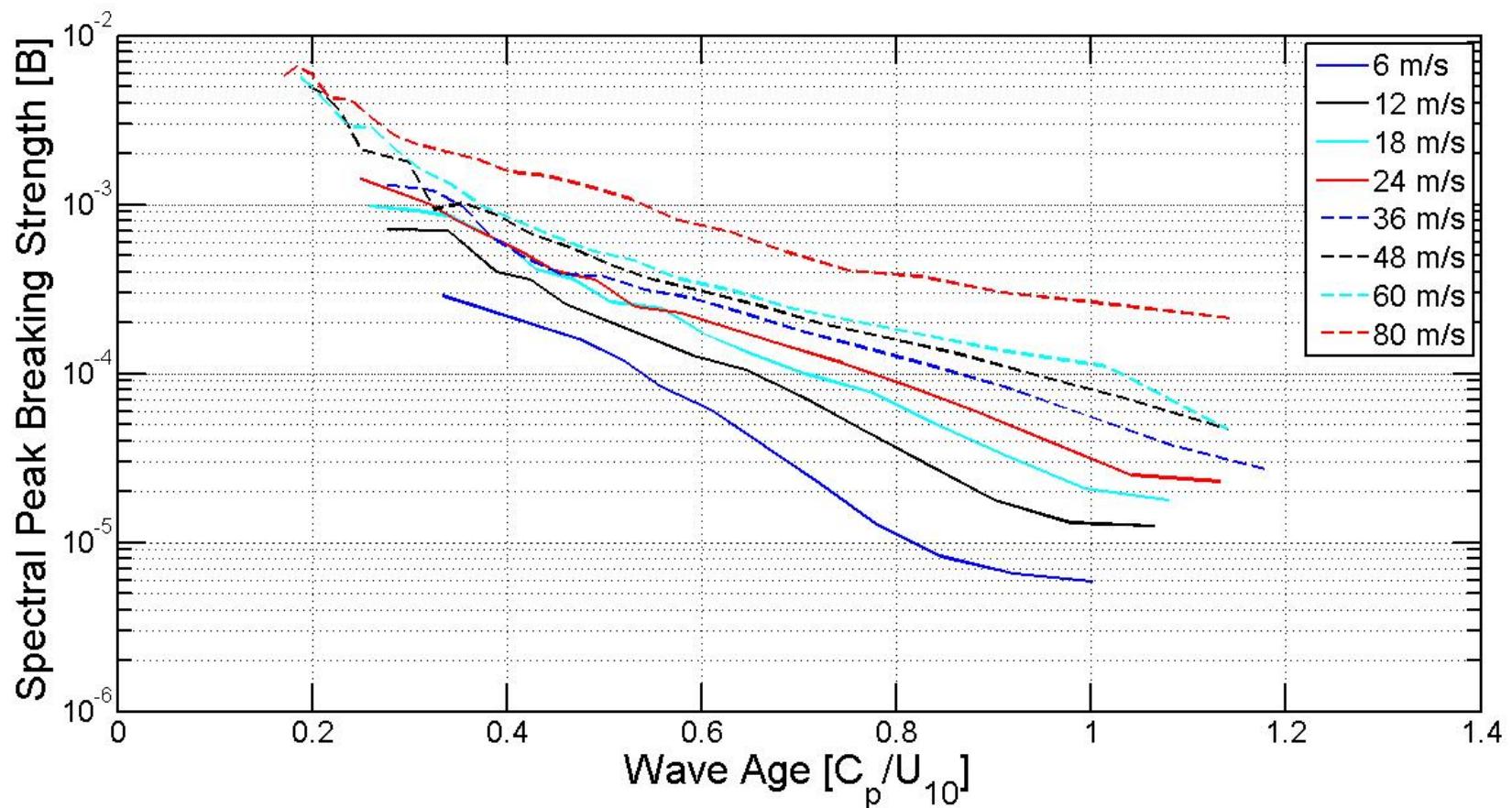


# Model Forecast: Breaking Probability .vs. Wave Age for multiple wind speeds 6 m/s to 80 m/s.

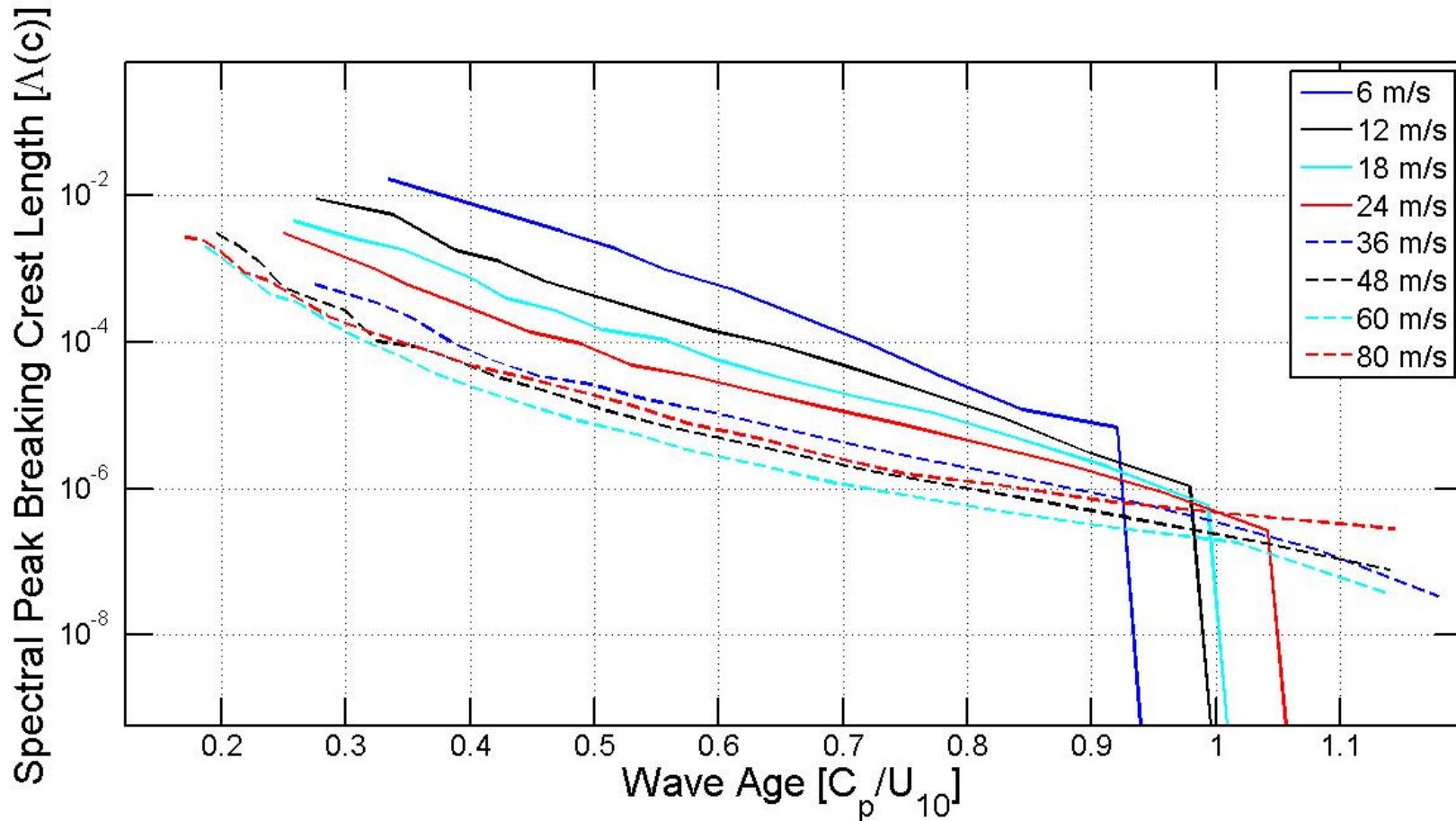


# Model Forecast: Breaking Strength .vs. Wave Age

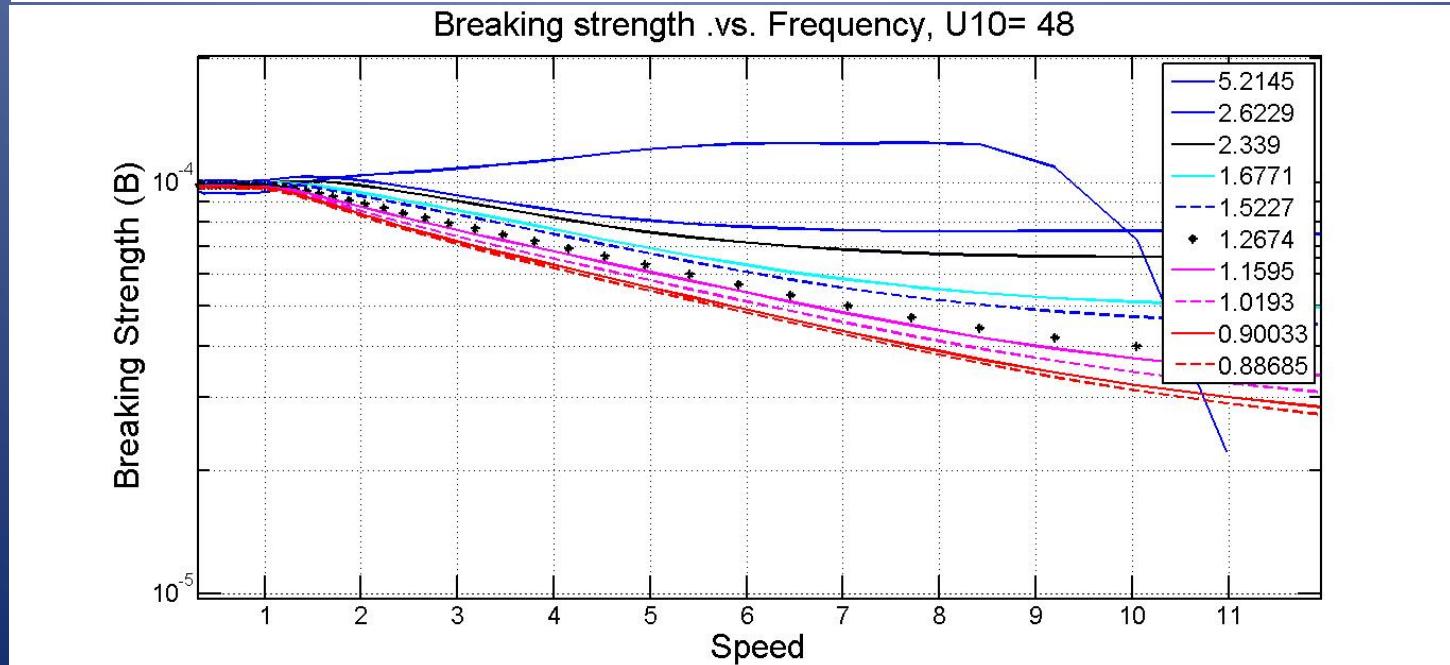
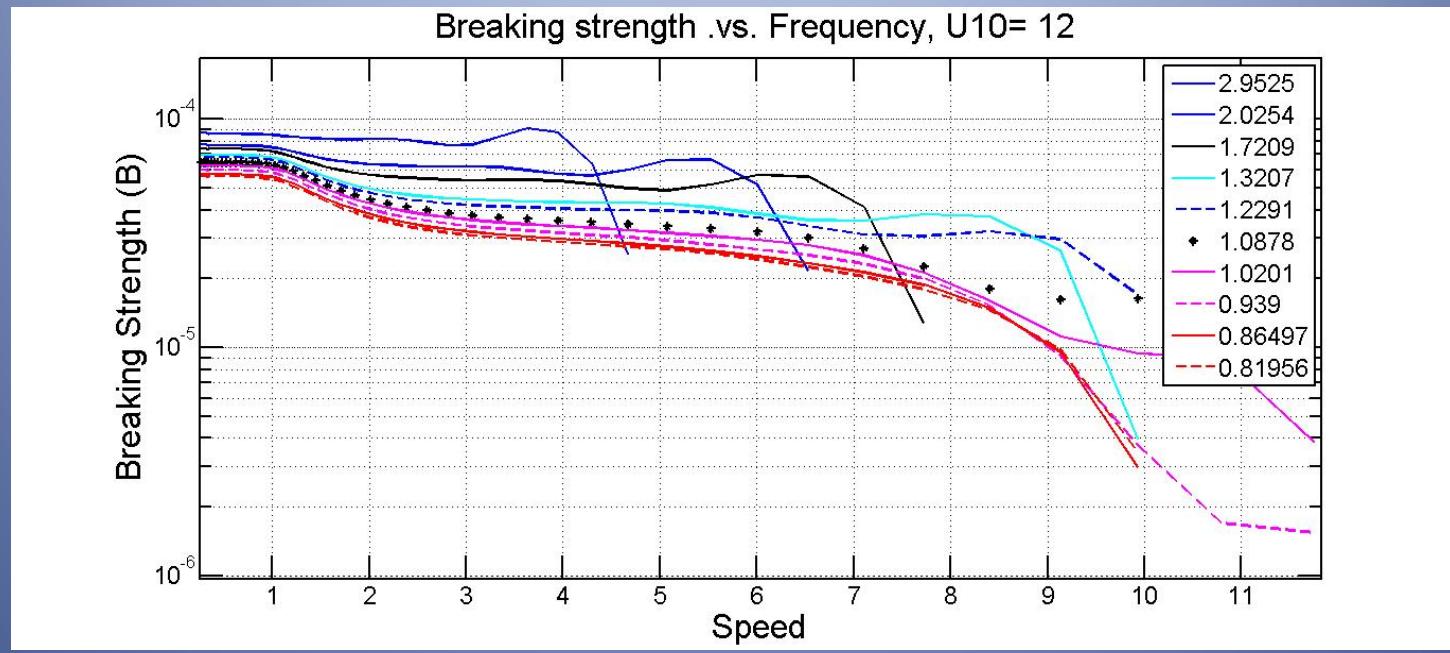
## For multiple wind speeds 6 ms/ to 80 m/s.



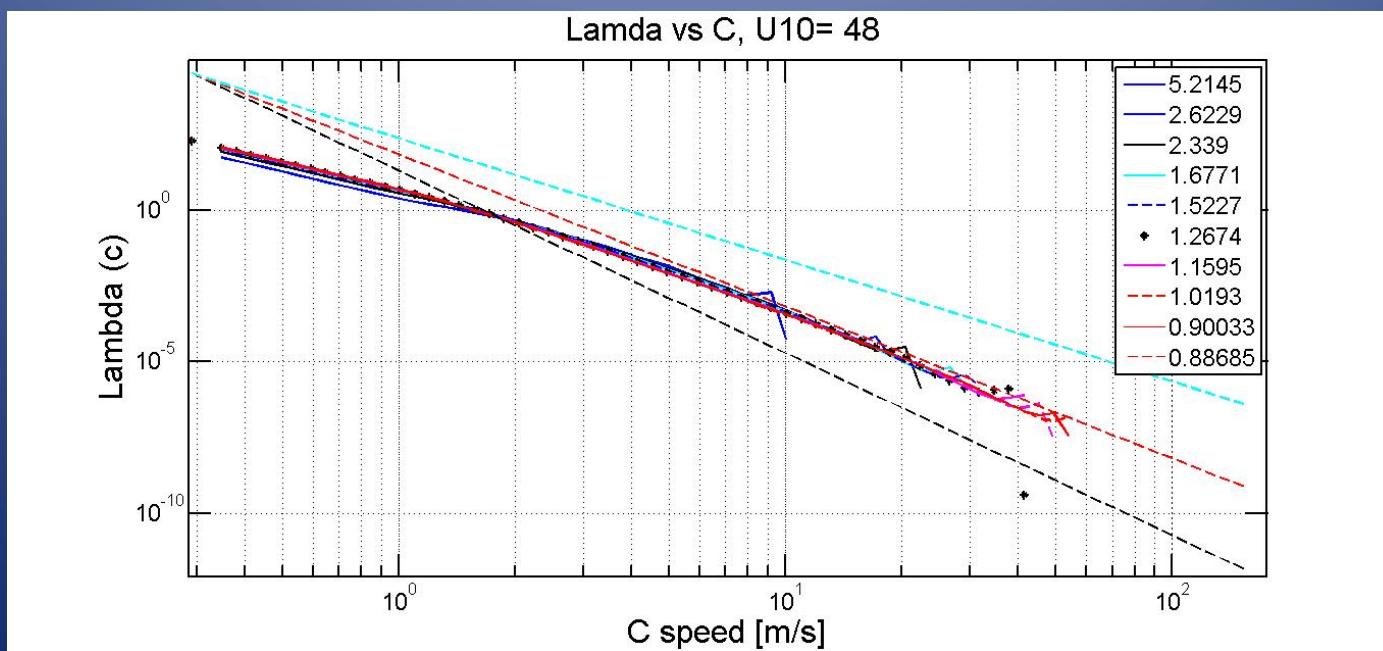
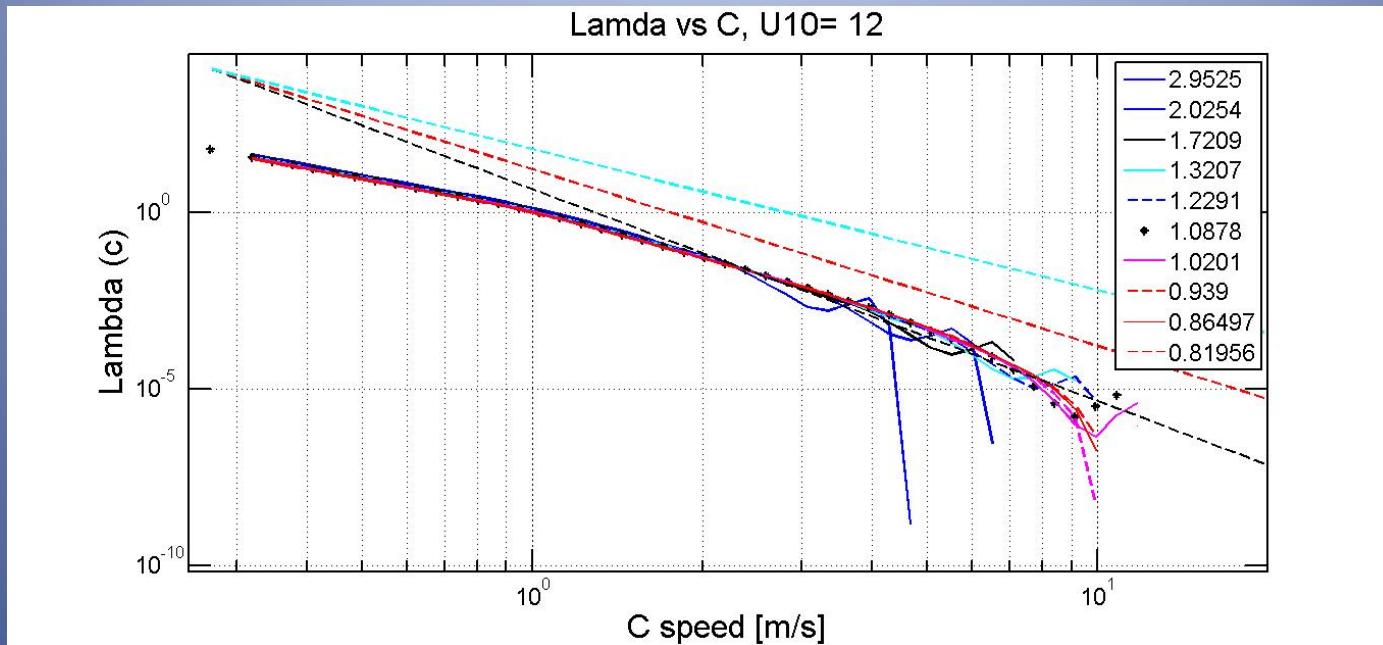
# Model Forecast: Lambda(c) .vs. Wave Age. For multiple wind speeds 6 ms/ to 80 m/s.

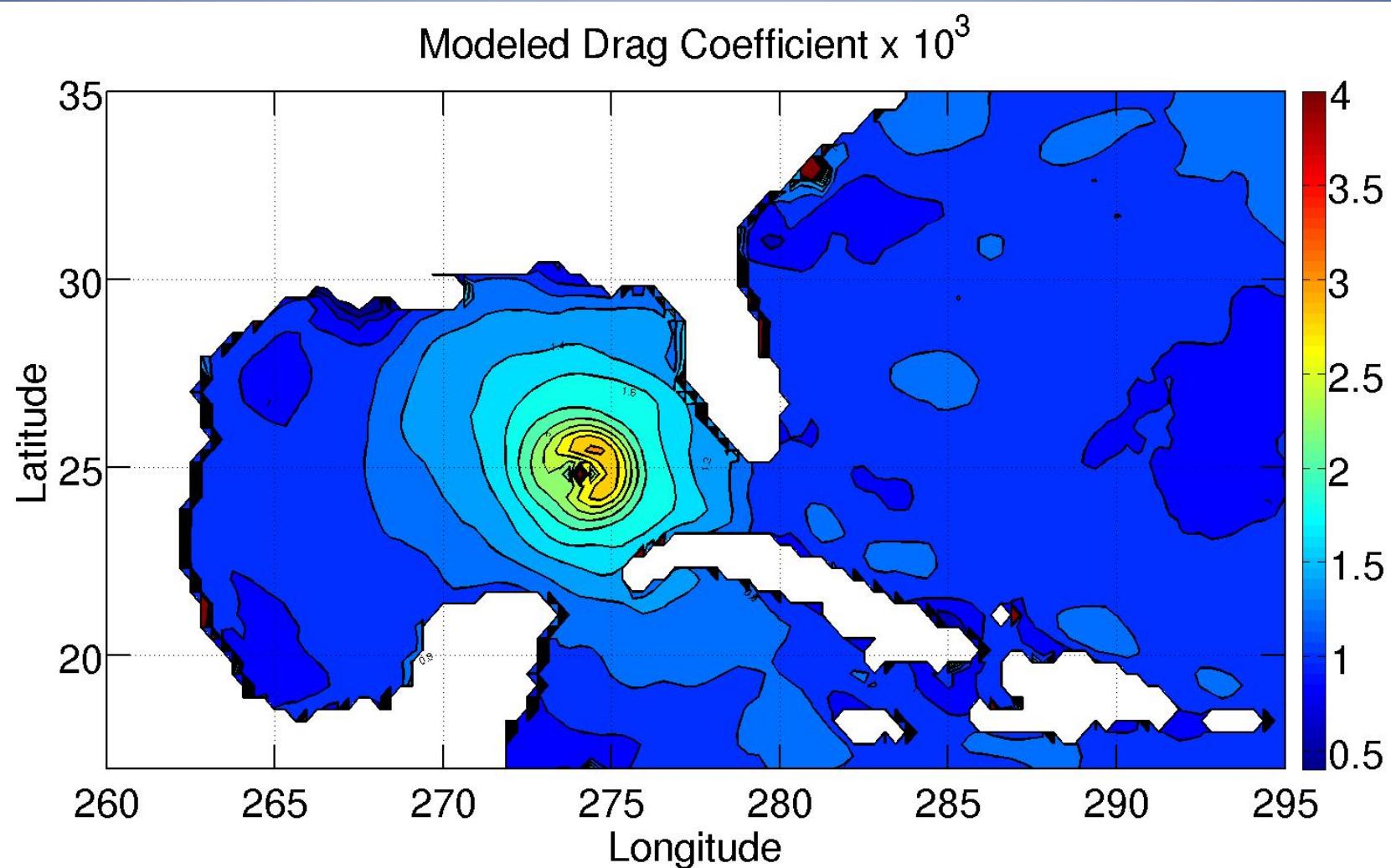


# Model Forecast: Breaking Strength [b] .vs. Speed for 12m/s & 48 m/s.

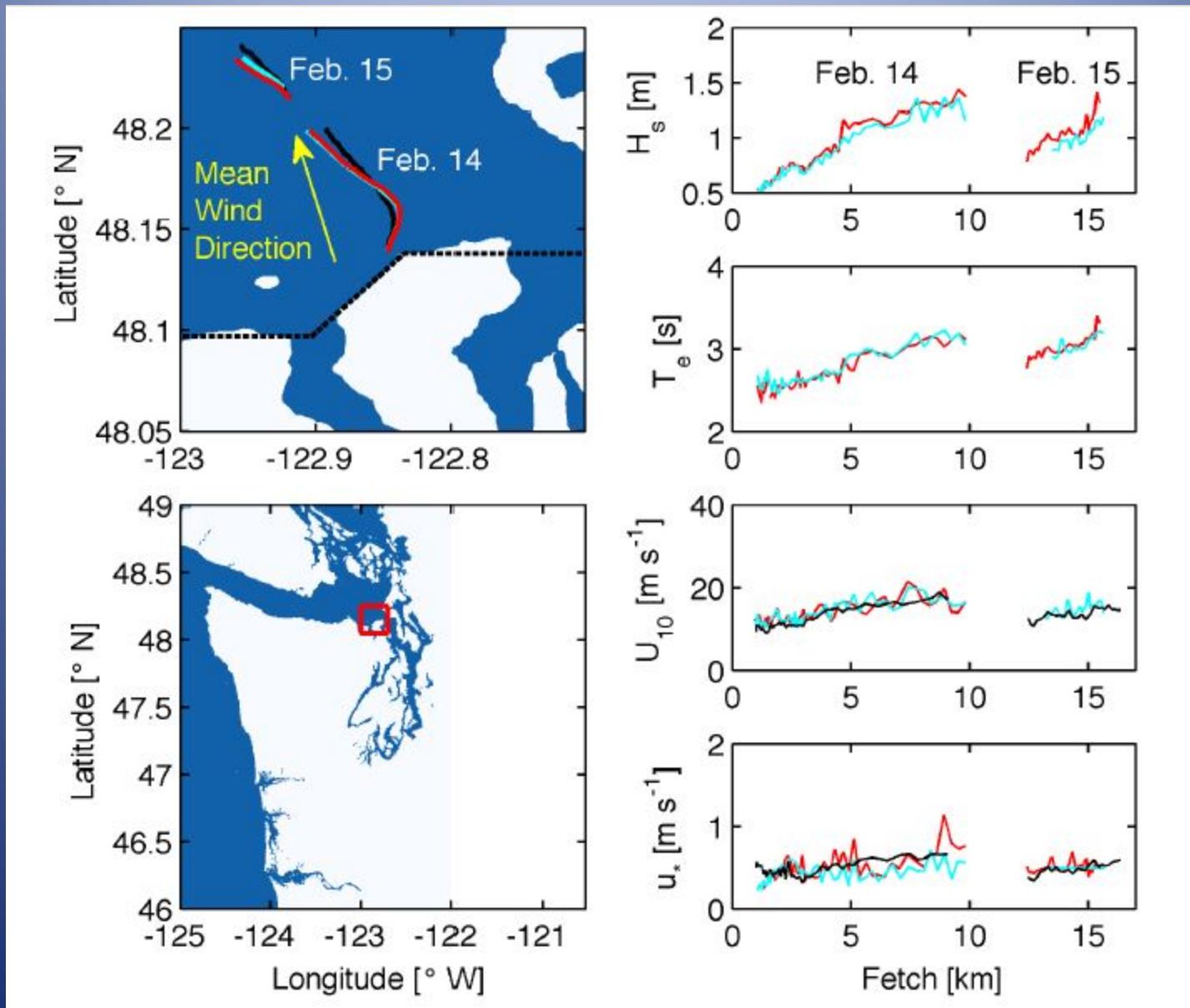


# Model Forecast: Lambda(c) .vs. Speed for 12 m/s & 48m/s.

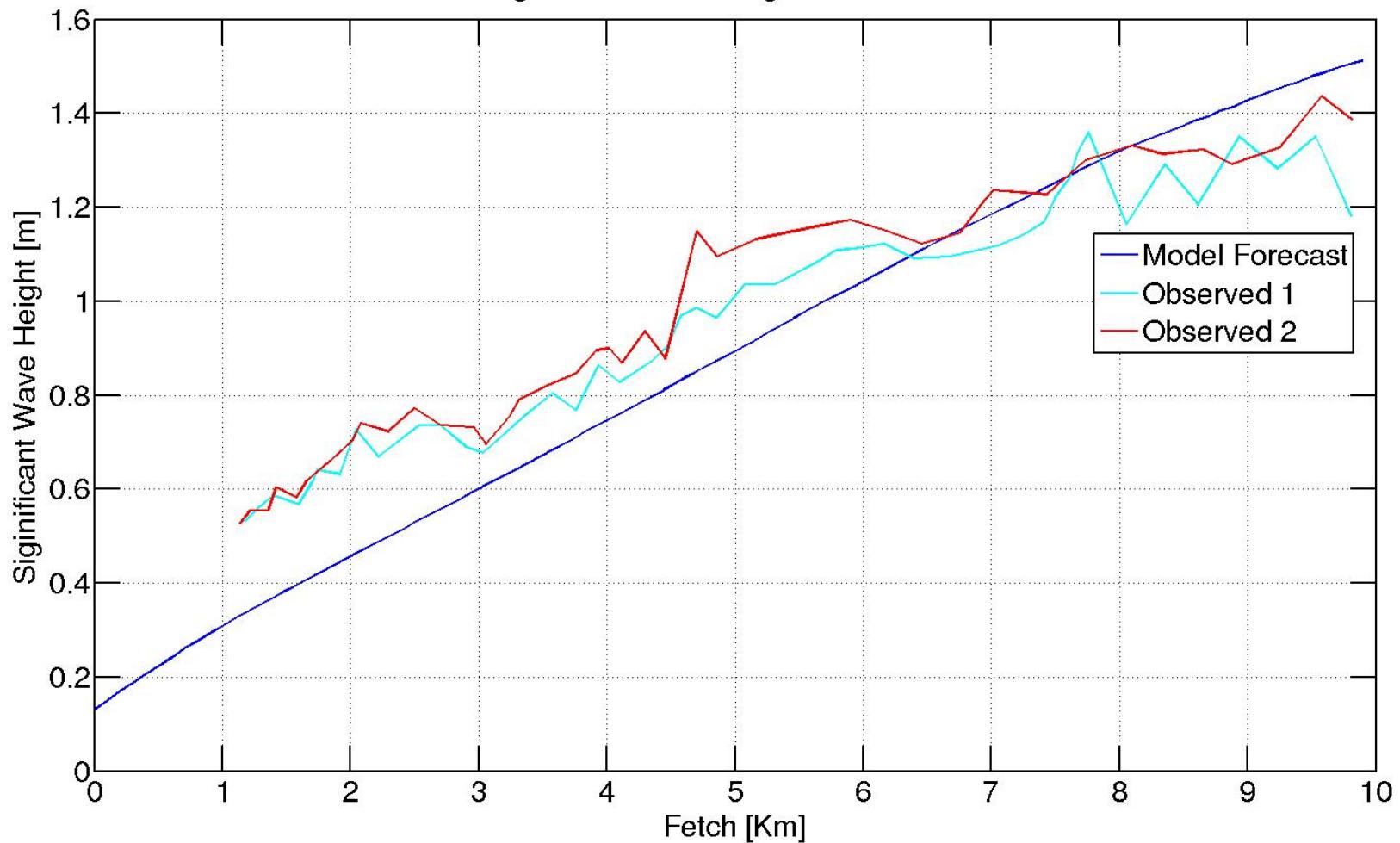


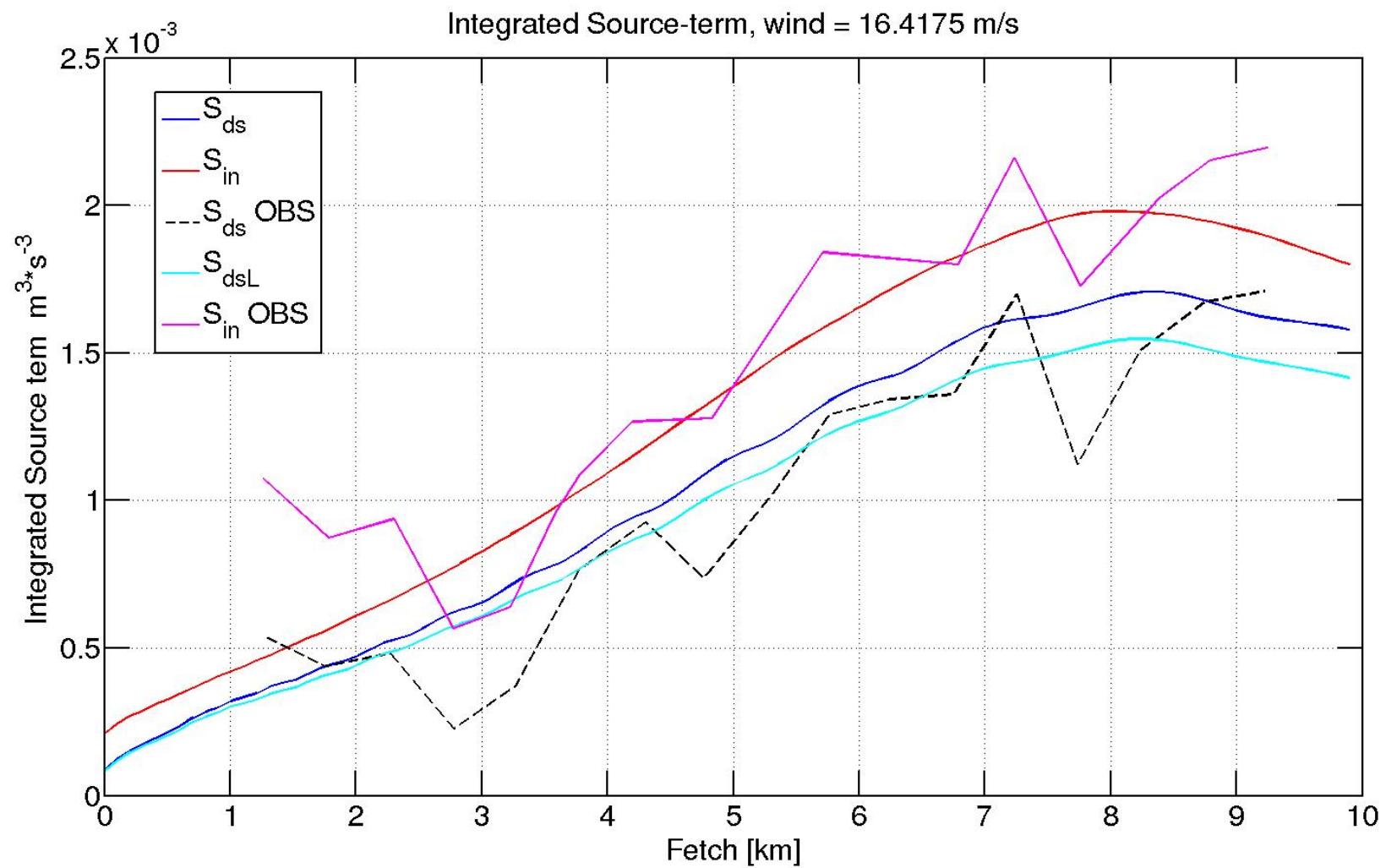


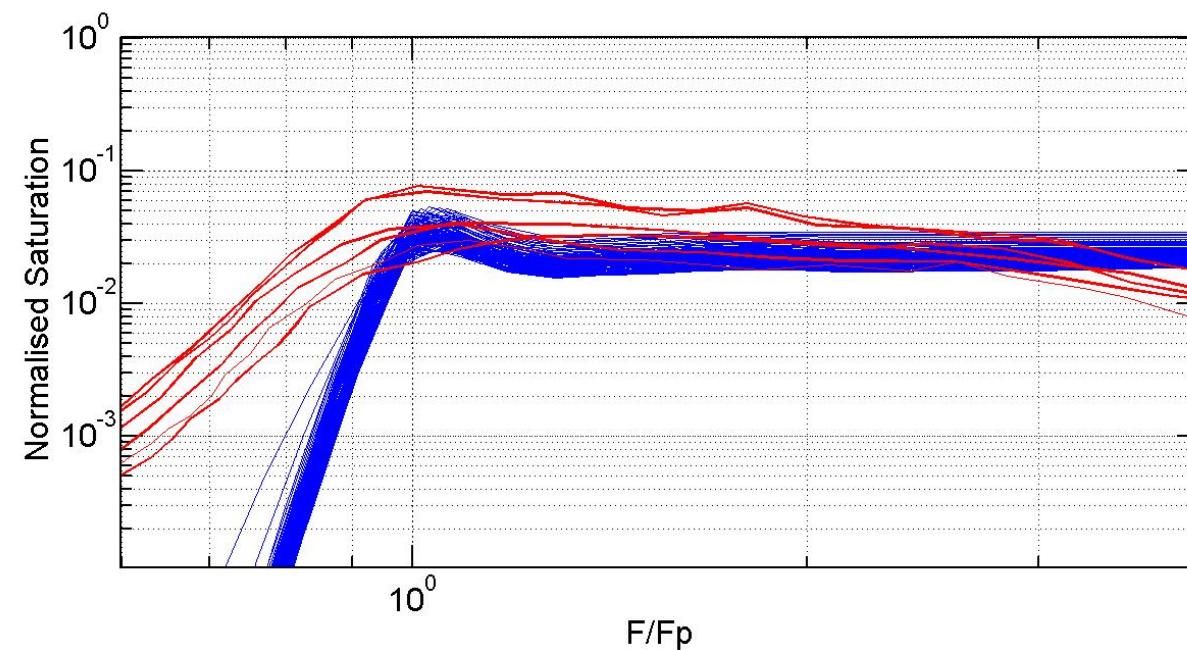
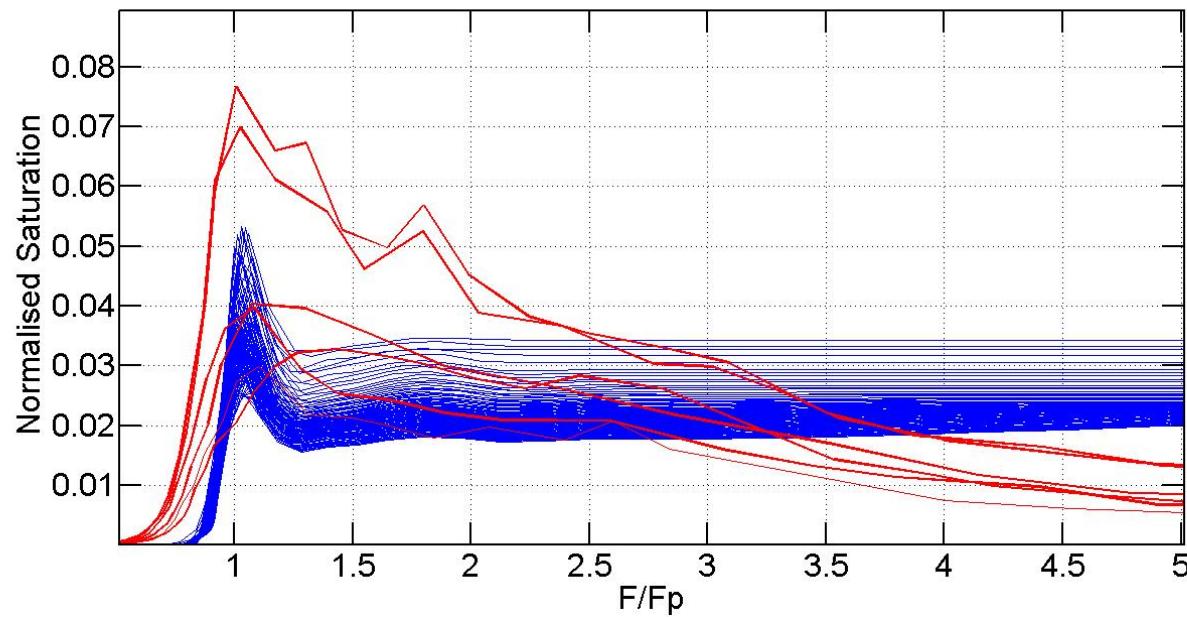
# Strait of Juan de Fuca: Experiment Conditions

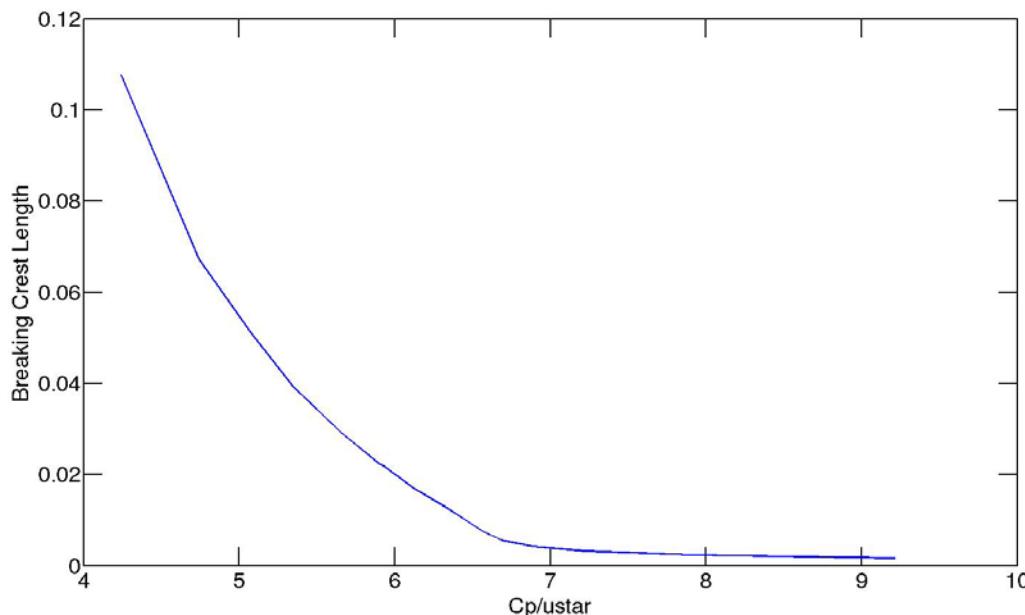
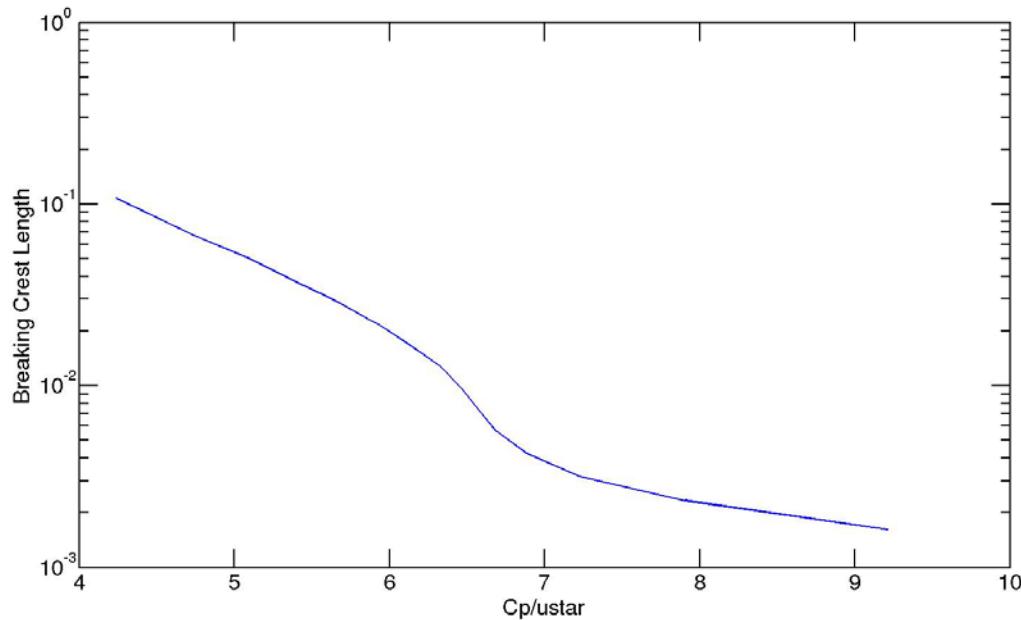


### Significant Wave Height versus Fetch









## Concluding Remarks

- our framework provides predictions of wave breaking properties (breaking probability, breaking crest length spectral density per unit area and breaking strength) using standard spectral wave models.
- it provides accurate predictions for the limited breaking data available in developing and mature wind seas
- further validation against data will be made as suitable new data sets become available.
- it has been added to existing spectral wave forecasting models. Upgrading the form of the DIA is desirable.

# We cant Save Everyone !

